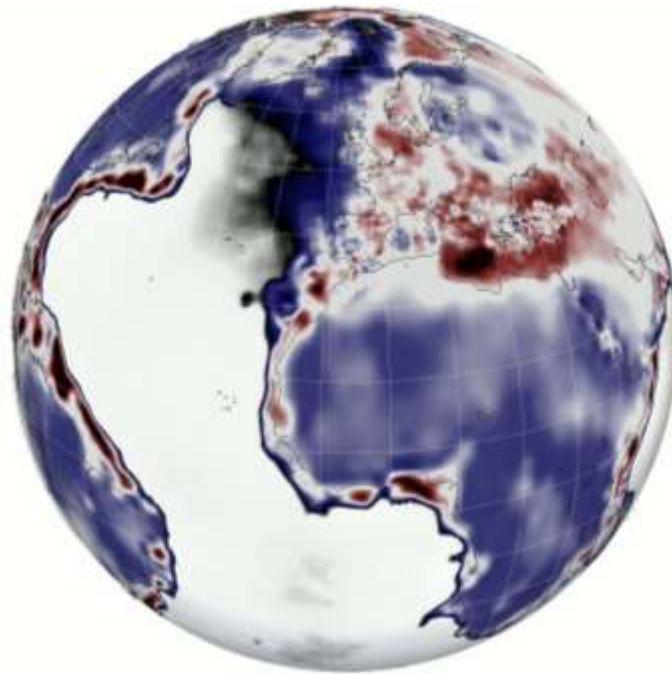


DEVELOPMENTS IN FULL-WAVEFORM INVERSION

Andreas Fichtner

Michael Afanasiev, Dirk-Philip van Herwaarden, Laura Cobden, Yesim Cubuk, Lion Krischer, Florian Rickers, Saulé Simuté, Jeannot Trampert, Tristan van Leeuwen, Antonio Villaseñor, *and many others ...*



v_{sv} at 15 km depth [km/s]
2.7 crust 4.0 mantle 4.8

“Full-waveform inversion (FWI) has emerged as the *final and ultimate solution* to the Earth resolution and imaging objective.”

Announcement of the 2013 SEG workshop on FWI

“Full-waveform inversion (FWI) has emerged as the *final and ultimate solution* to the Earth resolution and imaging objective.”

Announcement of the 2013 SEG workshop on FWI

- We do not know the Earth very well ...
... and are still looking for *the* method to solve all our problems.
- FWI seems poorly understood outside a small group of people.
Believe in miracles without seeing the limitations.
- Not generally understood that there are generally no ultimate solutions.
All methods have range of applicability.

OUTLINE

1. Full-waveform inversion in a nutshell

- From ray tomography to full-waveform inversion
- A synthetic illustration of the main benefit

2. Real-world examples

- The Japanese islands: Recovering extremely low mantle velocities
- The Western Mediterranean: Crust/mantle resolution and uncertainty analysis

OUTLINE

1. Full-waveform inversion in a nutshell

- From ray tomography to full-waveform inversion
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2. Real-world examples

- The Japanese islands: Recovering extremely low mantle velocities
- The Western Mediterranean: Crust/mantle resolution and uncertainty analysis

3. Big challenges

- Are our computers big enough, or will they be in the near future?
- The data flood
- Our mode of operation

4. The Collaborative Seismic Earth Model

- Philosophy and technical implementation
- Generation 1

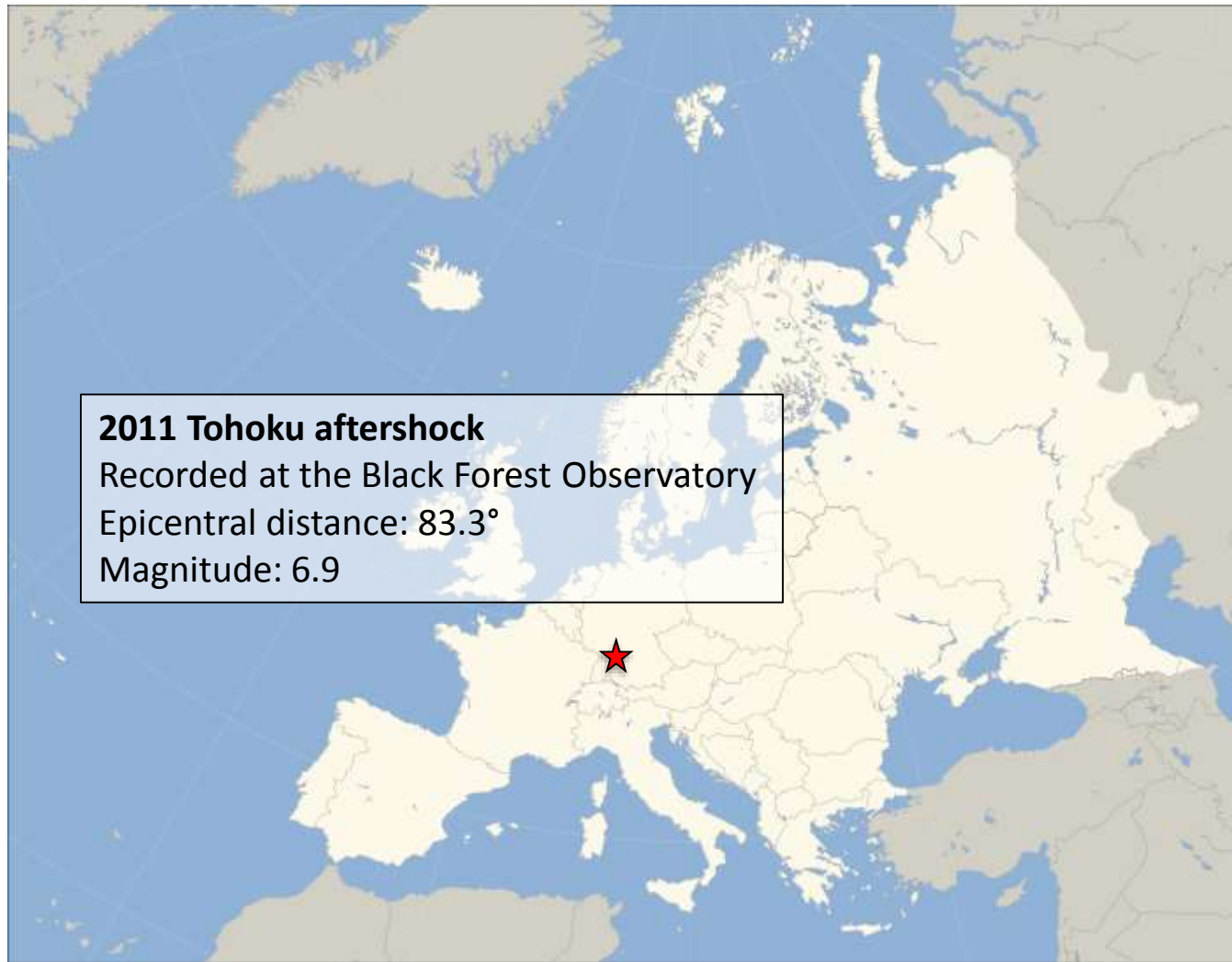
5. Discussion and Conclusions

FULL-WAVEFORM INVERSION IN A NUTSHELL

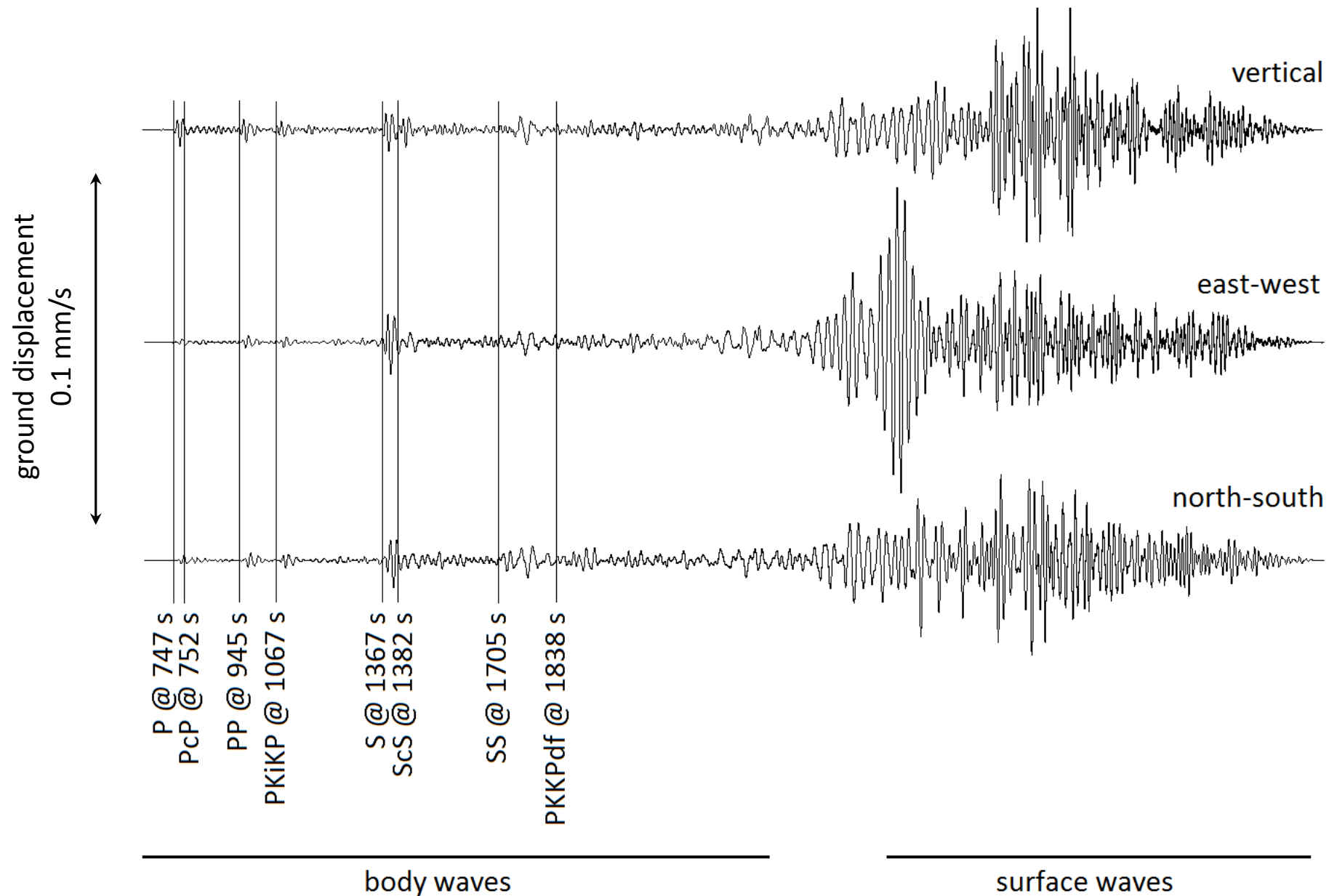
Exploiting complete waveforms for the benefit of improved resolution



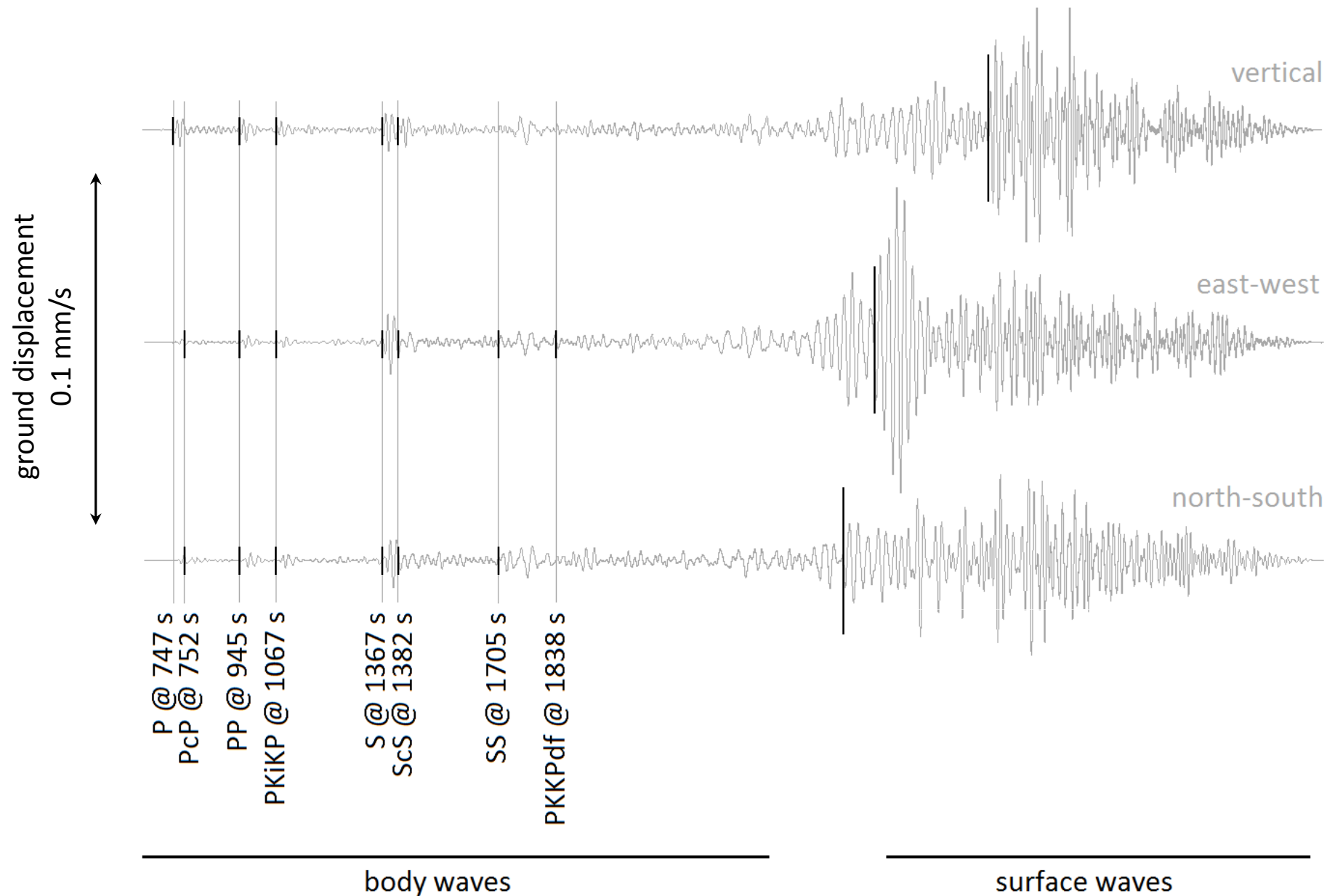
Numerical simulation [using SALVUS; Afanasiev et al., 2018] of a Tohoku aftershock [M 6.9]



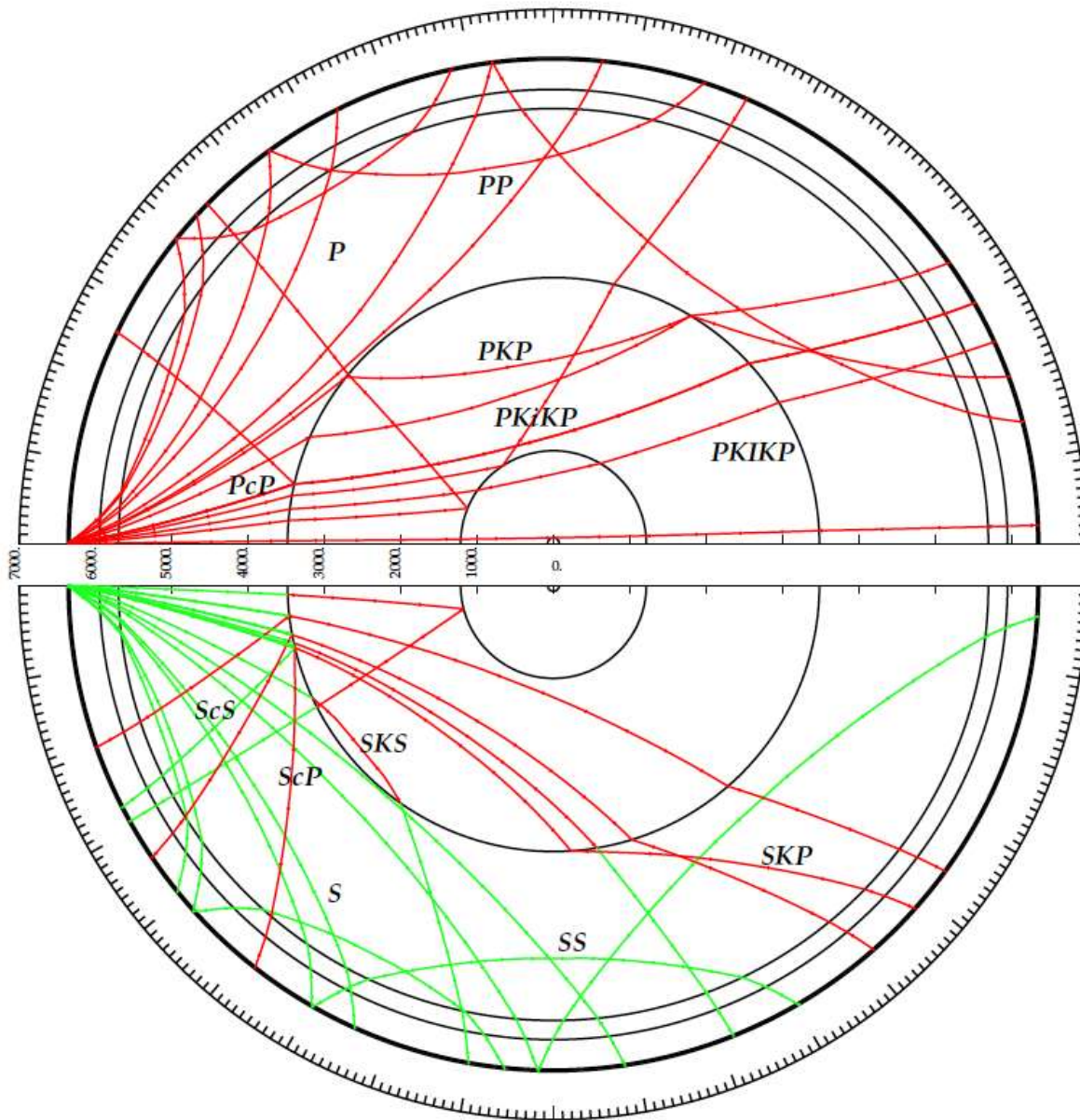
SEISMOGRAMS



1: TRAVELTIME RAY TOMOGRAPHY

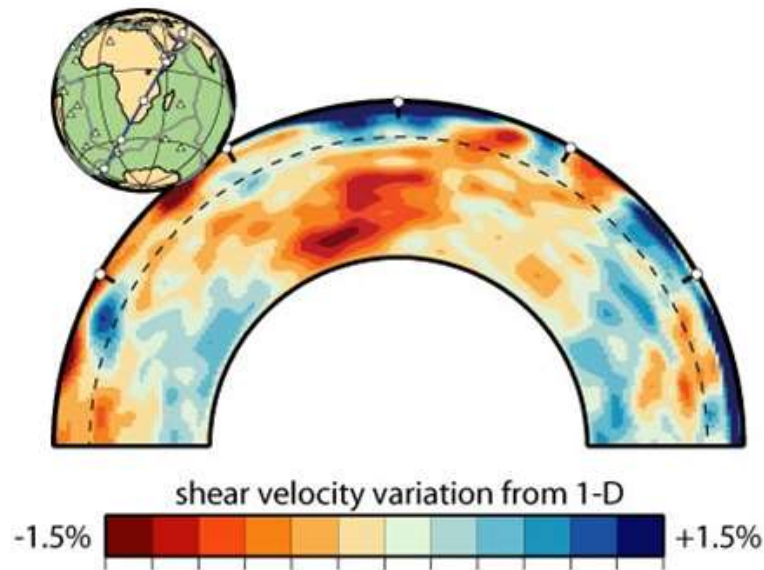


1: TRAVELTIME RAY TOMOGRAPHY

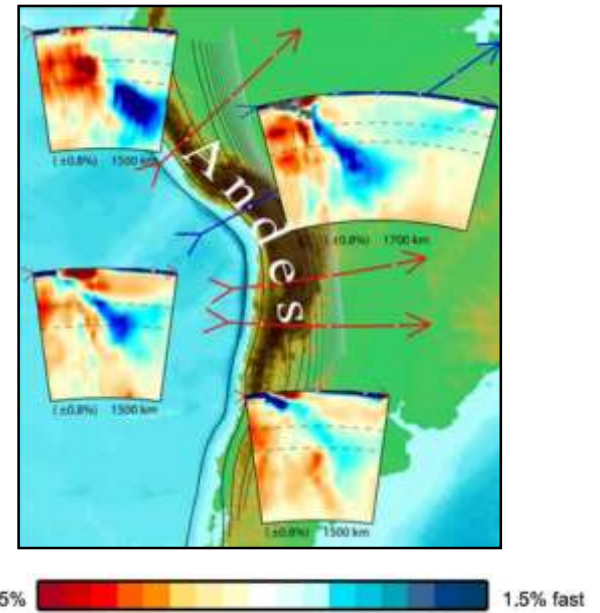


ak135 traveltime tables, Kennett et al., 2005

1: TRAVELTIME RAY TOMOGRAPHY



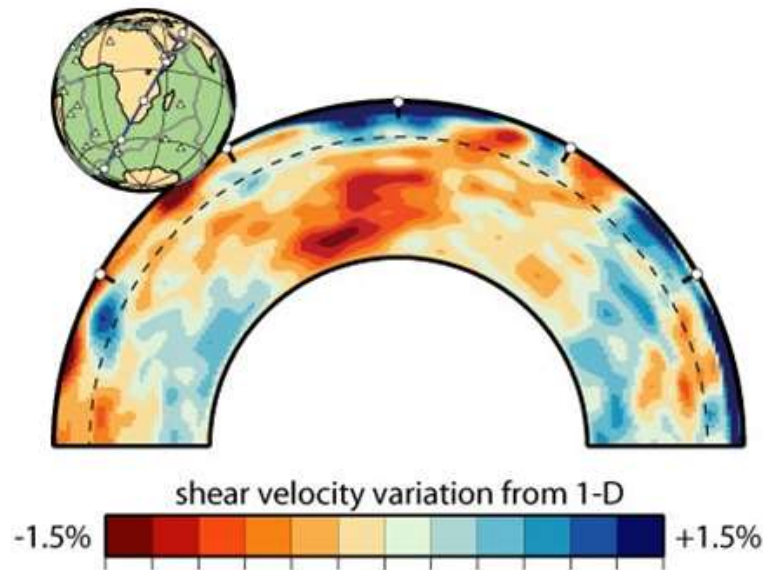
Ritsema et al., 2010



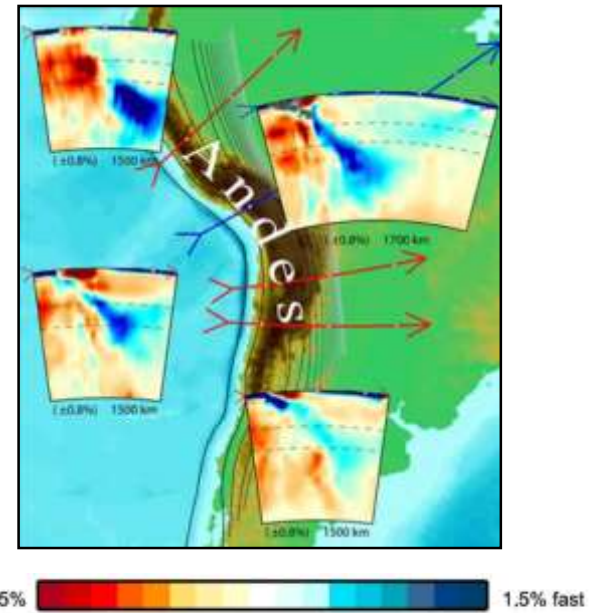
Obayashi et al., 1997

- The majority [by far] of all tomographies are traveltime ray tomographies.
- First applications: Aki et al. [1977], Dziewonski et al. [1977].
- Very well established.

1: TRAVELTIME RAY TOMOGRAPHY



Ritsema et al., 2010



Obayashi et al., 1997

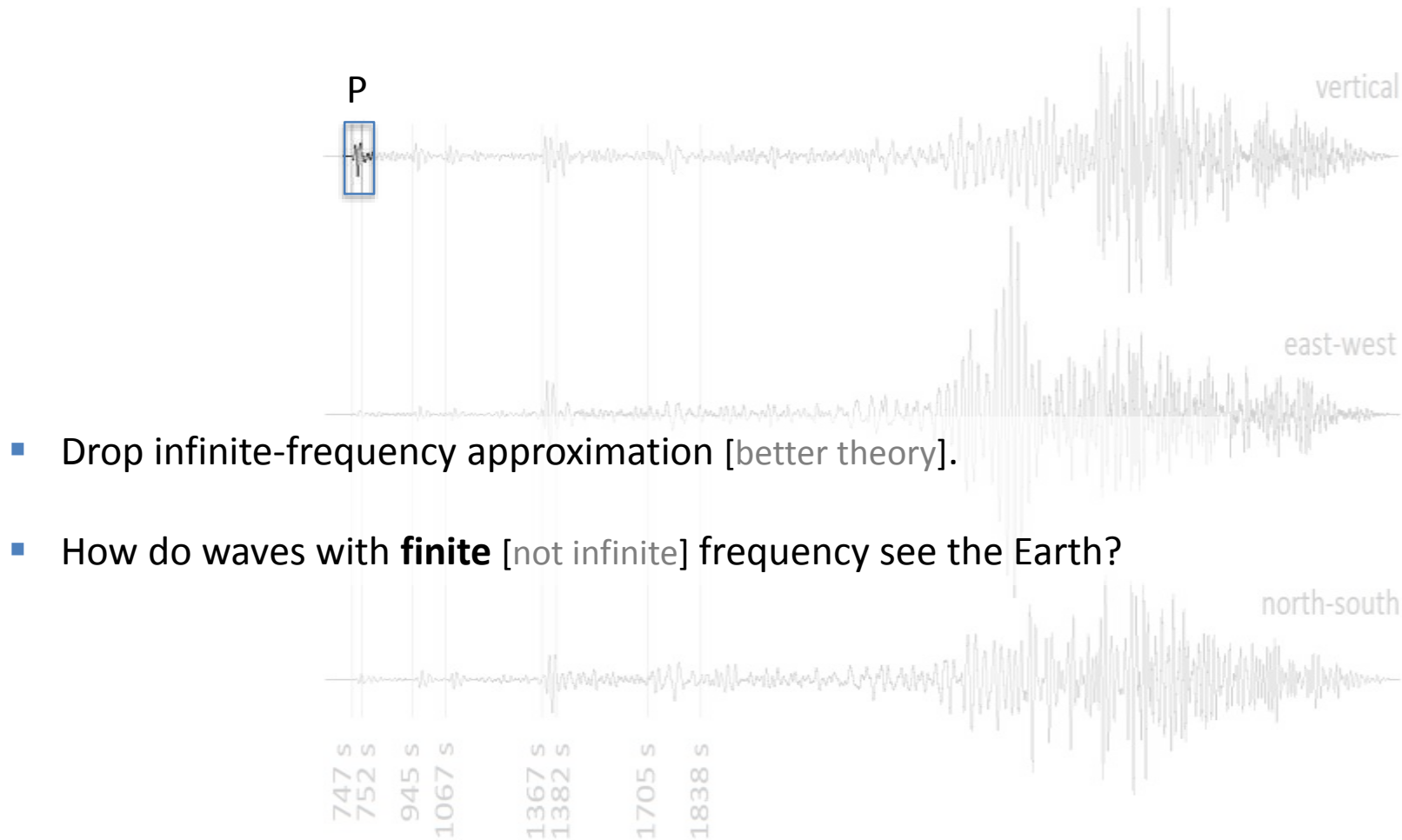
Pros

- Relatively simple theory.
- Computationally inexpensive.
- Possibility to incorporate a very large number of measurements.

Cons

- Ray theory is an **infinite-frequency** approximation for **smooth** media.
- Any information contained in the waveform details is ignored.

2: FINITE-FREQUENCY TOMOGRAPHY

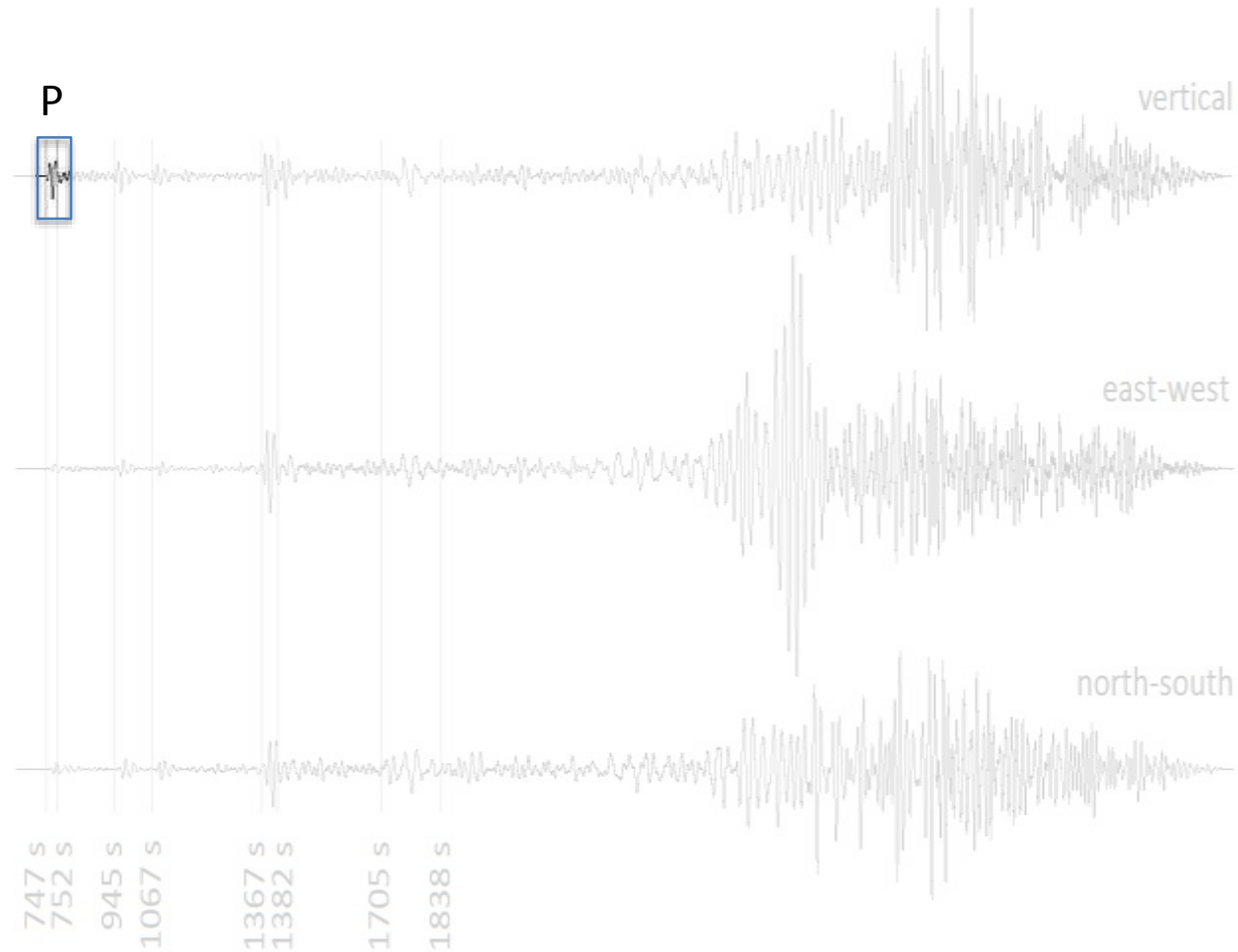
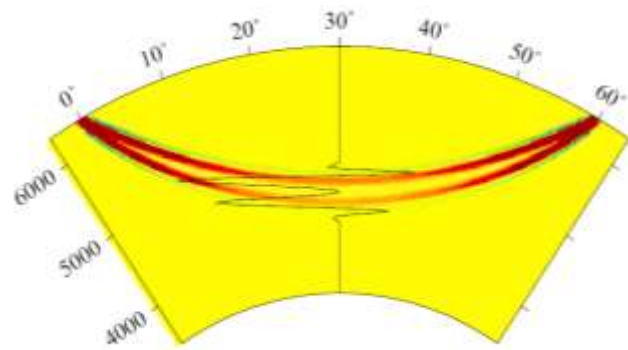


2: FINITE-FREQUENCY TOMOGRAPHY

Wave type: P wave

Parameter: P velocity

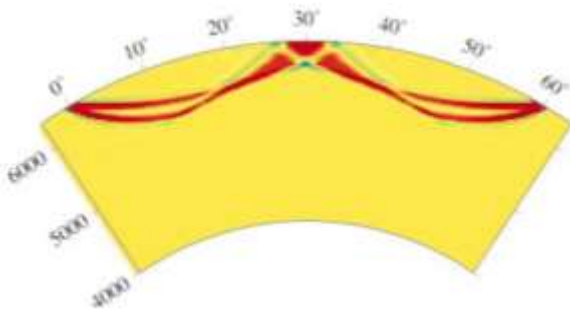
Period: 10 s



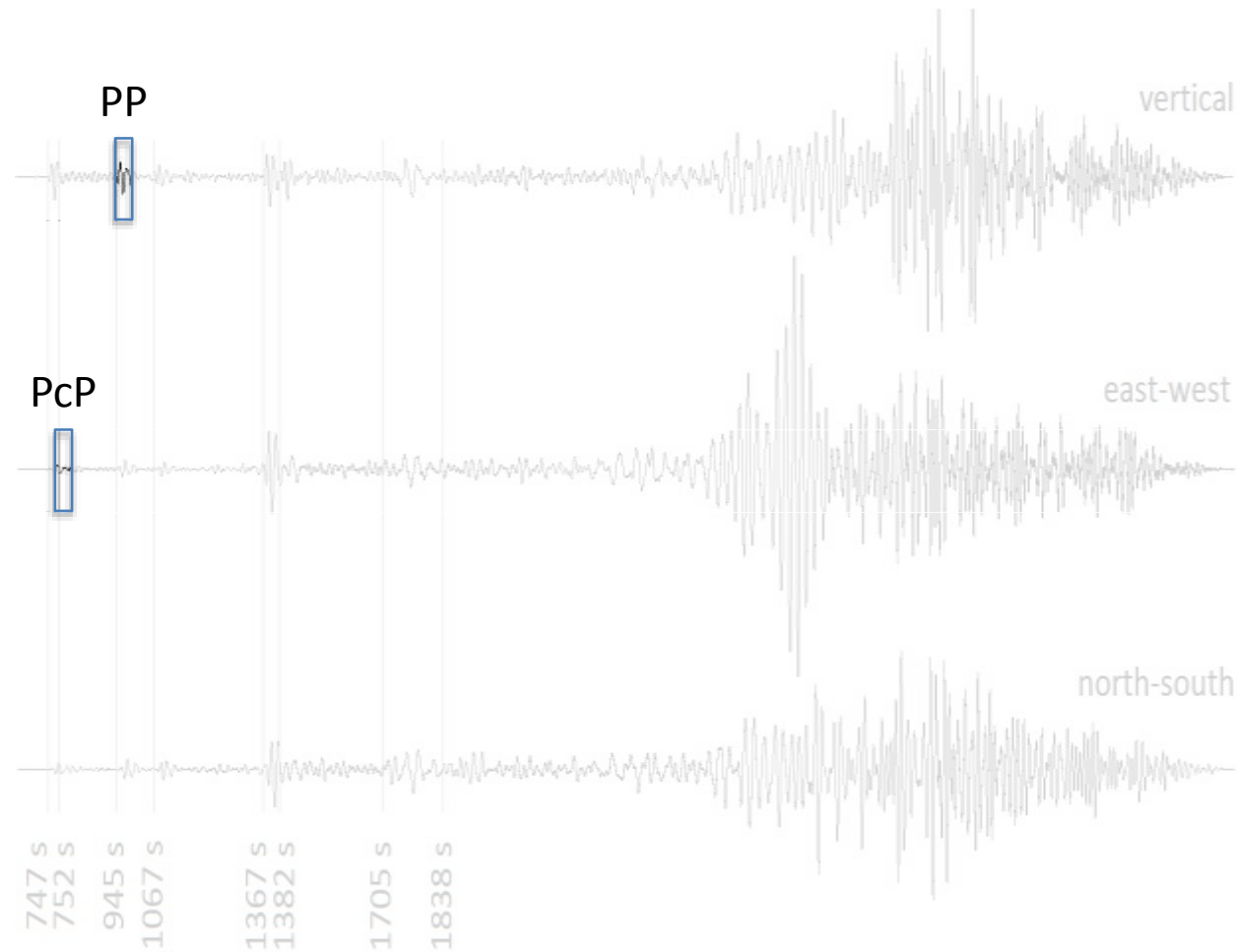
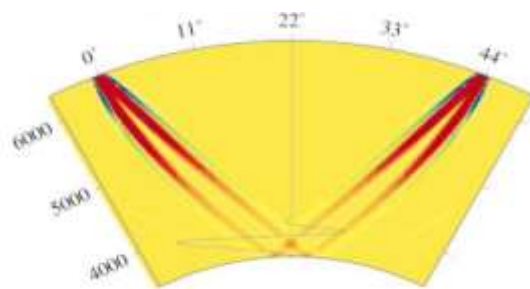
Hung et al. 2000

2: FINITE-FREQUENCY TOMOGRAPHY

Wave type: PP wave
Parameter: P velocity
Period: 20 s



Wave type: PcP wave
Parameter: P velocity
Period: 20 s



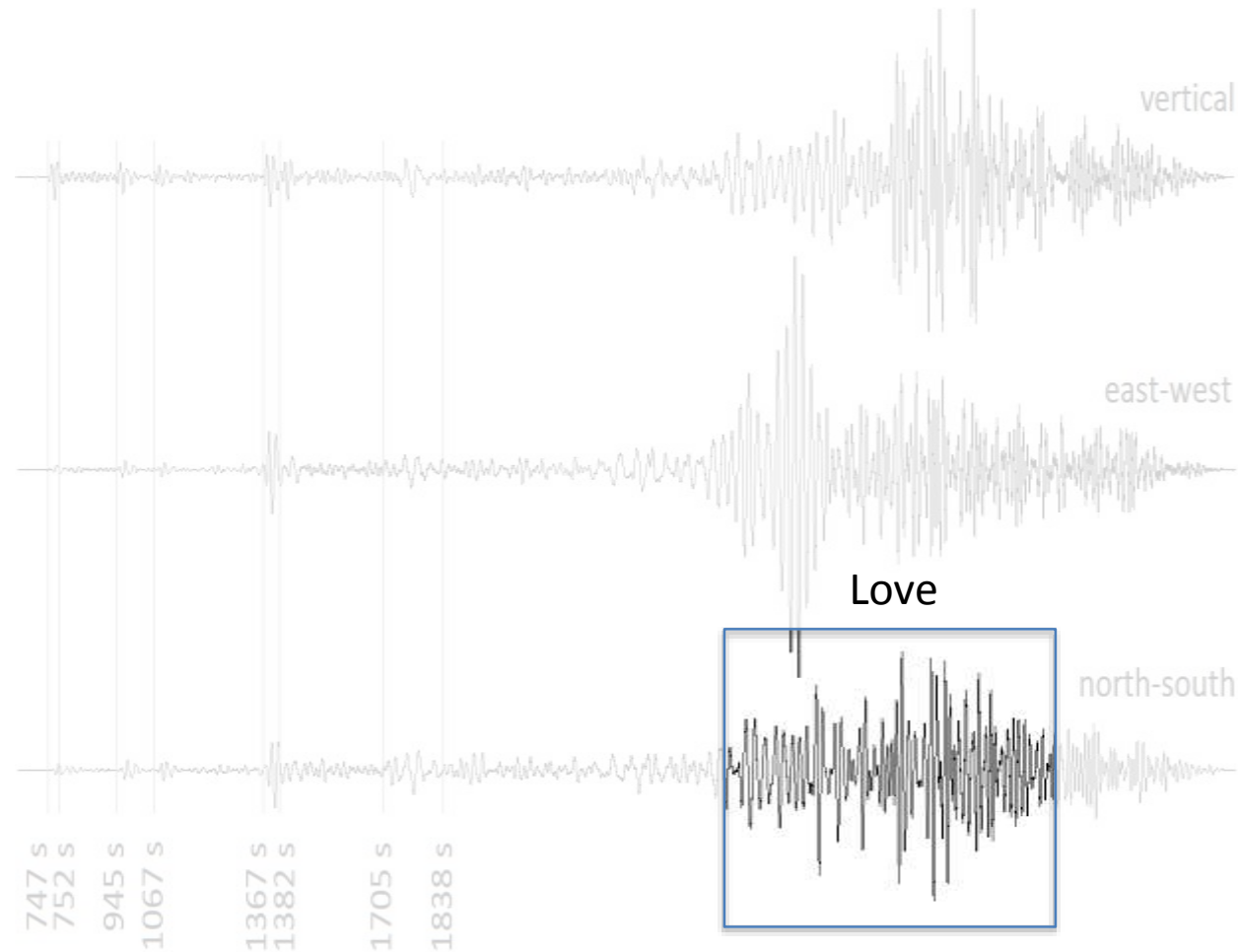
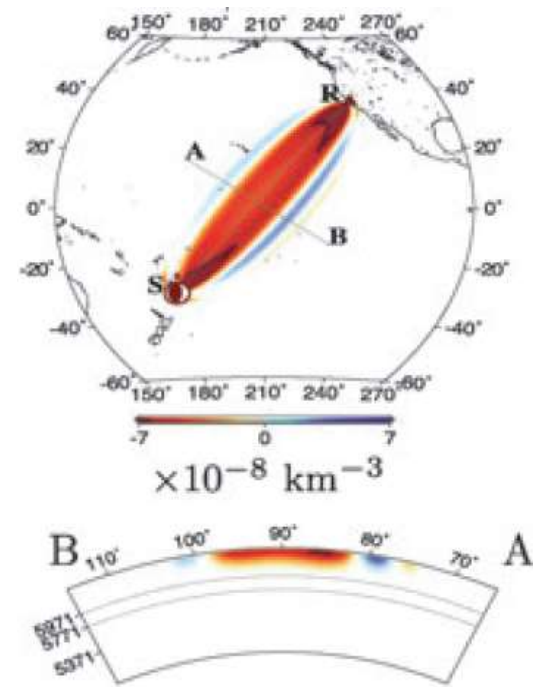
Hung et al. 2000

2: FINITE-FREQUENCY TOMOGRAPHY

Wave type: Single-mode Love wave

Parameter: S velocity

Period: 100 s



Zhou et al. 2004

2: FINITE-FREQUENCY TOMOGRAPHY

Wave type: Single-mode Love wave

Parameter: S velocity

Period: 100 s



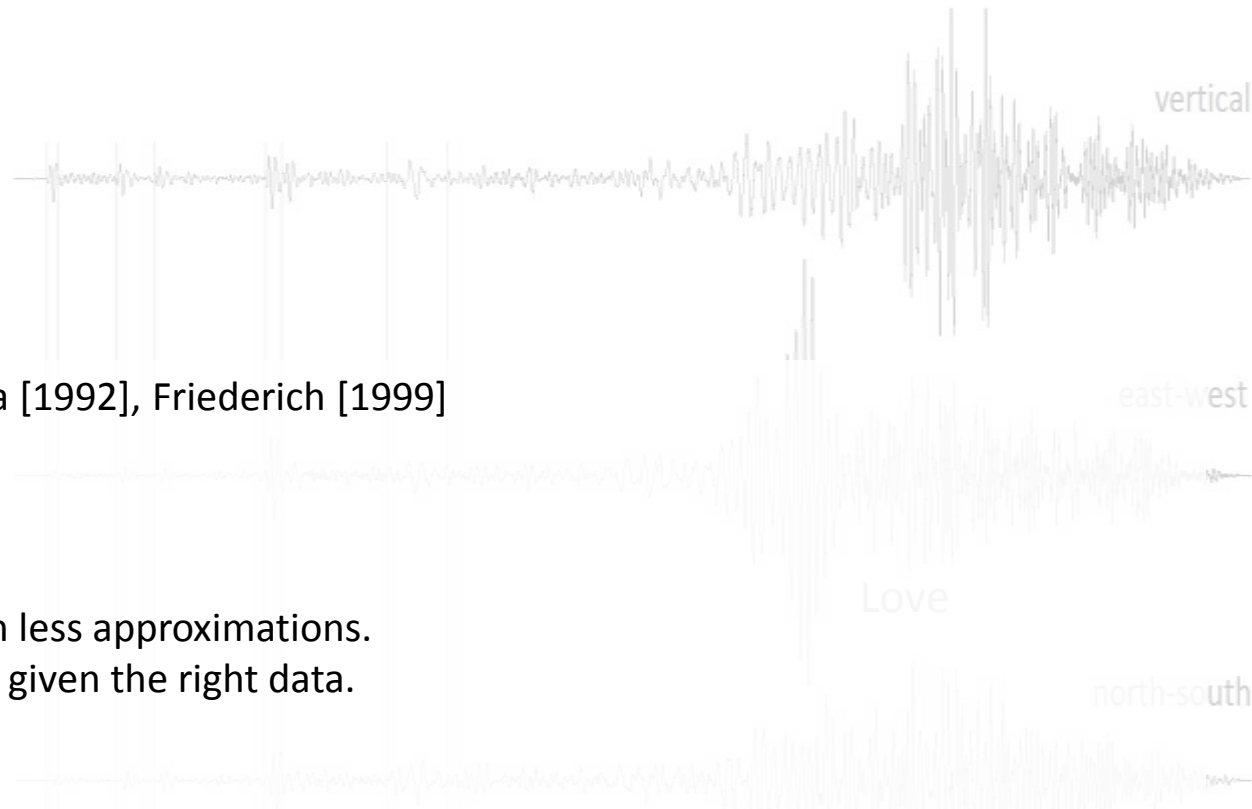
- First applications: Yomogida [1992], Friederich [1999]

Pros

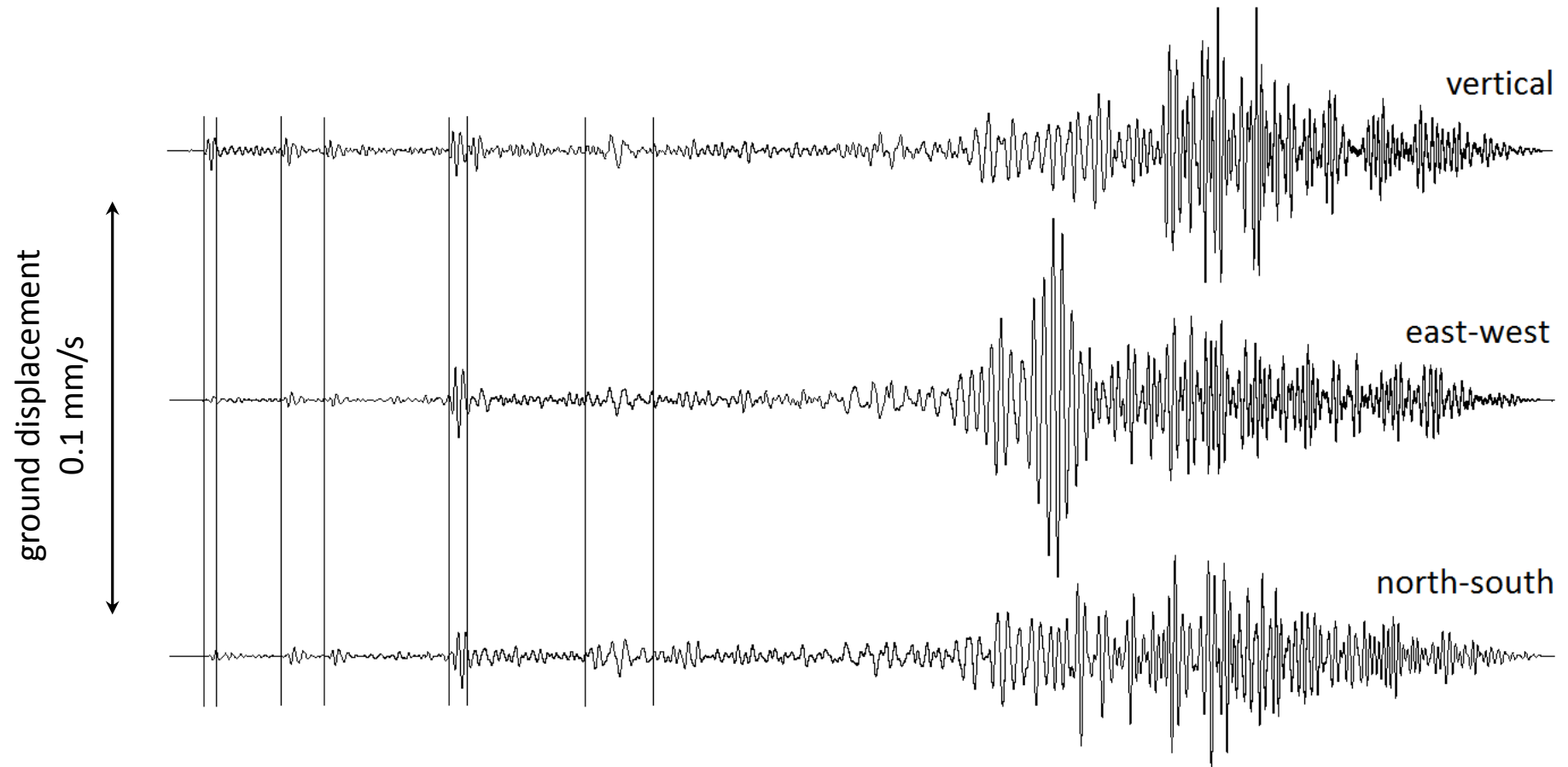
- More elaborate theory with less approximations.
- In theory better resolution, given the right data.

Cons

- Benefits are somewhat debated.
- Still ignores anything that is not a well-defined phase.
- Variants developed so far also require smooth media.



3: FULL-WAVEFORM INVERSION



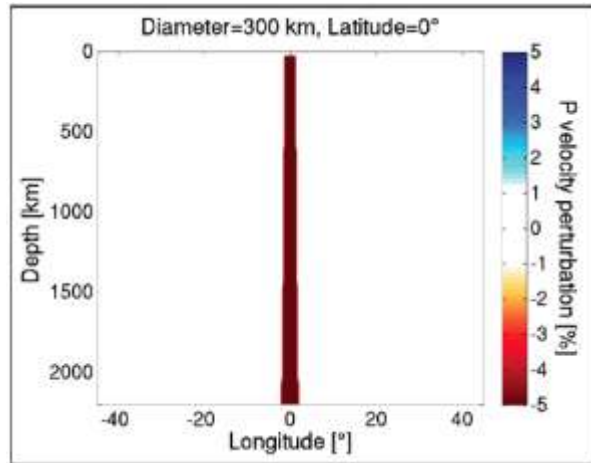
Solve the wave equation fully numerically for heterogeneous Earth [see Heiner Igel's talk].

- Avoid any significant modelling error [and related imaging artifacts].
- Use as much information as you can [every wiggle, if the noise permits to do so].

3: FULL-WAVEFORM INVERSION

SYNTHETIC EXAMPLE

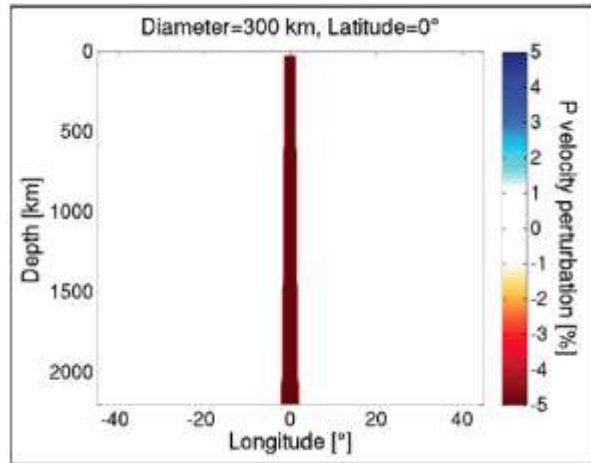
Input model – an idealised plume



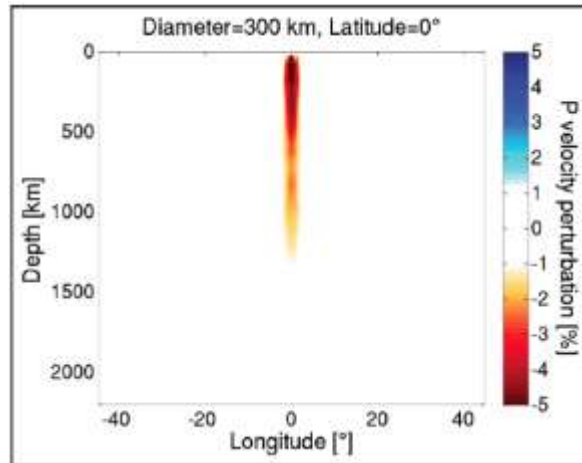
3: FULL-WAVEFORM INVERSION

SYNTHETIC EXAMPLE

Input model – an idealised plume



reconstructed model
using only P wave traveltimes



Wavefront healing: direct wave forgets about the plume.

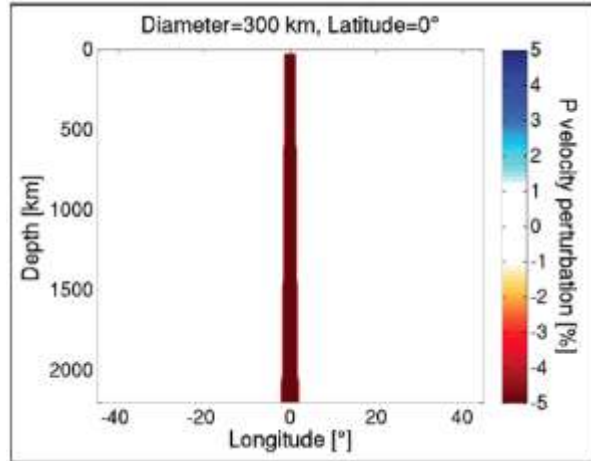
Deep plumes cannot be resolved with traveltime tomography

[e.g. Trembl 2006, Hwang 2011, Rickers 2012, Maguire et al. 2016].

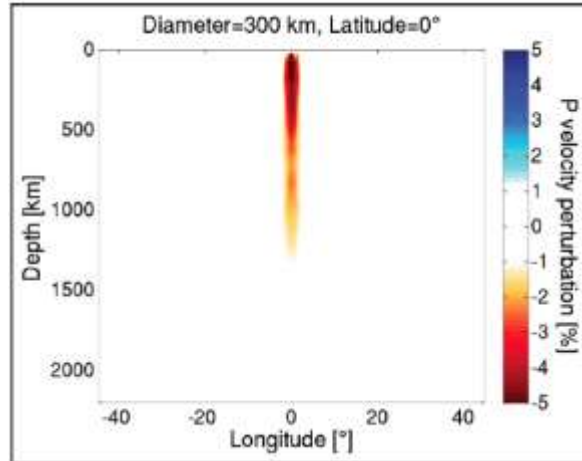
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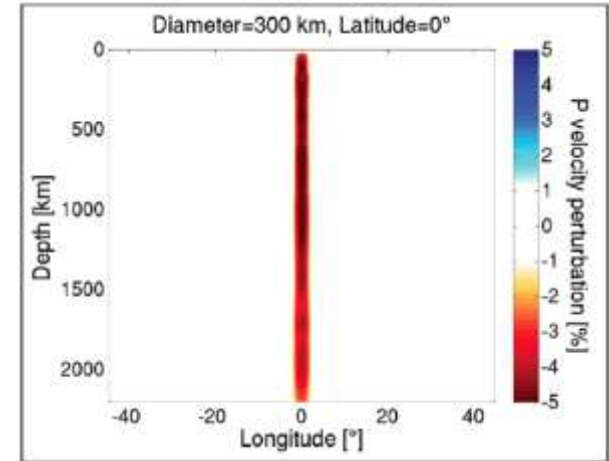
Input model – an idealised plume



reconstructed model
using only P wave traveltimes



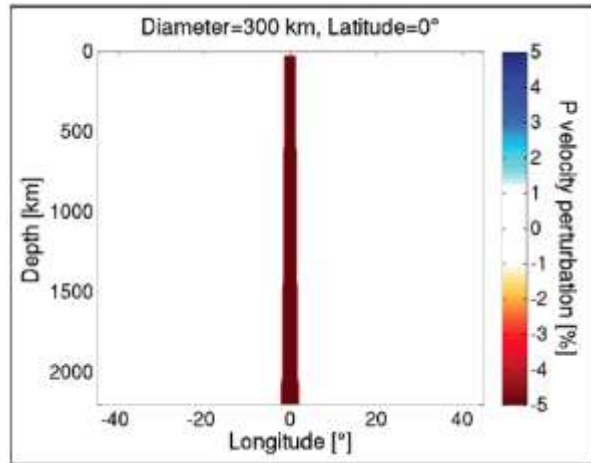
reconstructed model
complete seismograms



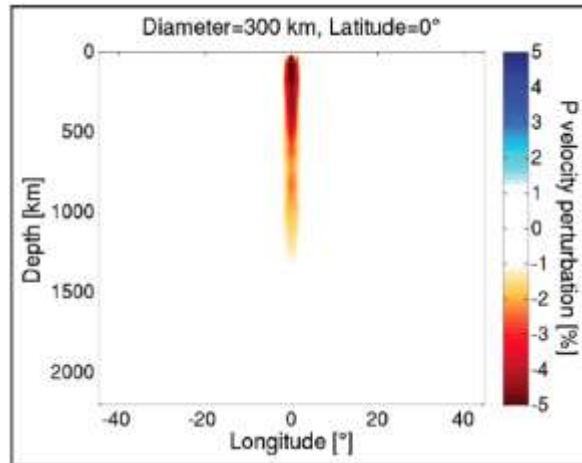
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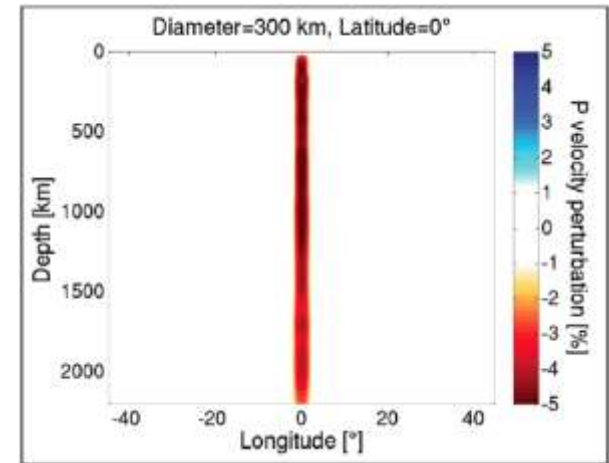
Input model – an idealised plume



reconstructed model
using only P wave traveltimes



reconstructed model
complete seismograms

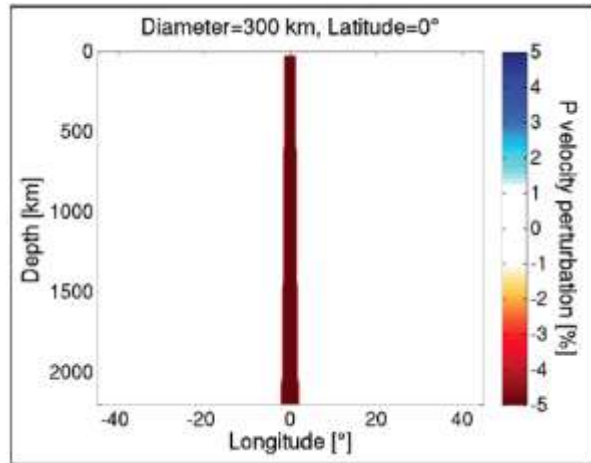


- First theoretical attempts: Bamberger, Chavent, Lailly [late 1970's]
- First applications in 2D: Crase, Igel, Tarantola [1990's]
- First applications in 3D: Chen, Tape, Fichtner [nearly 10 years ago]

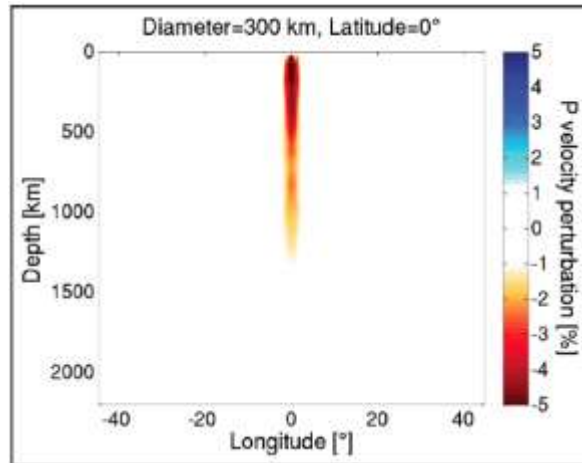
3: FULL-WAVEFORM INVERSION

SYNTHETIC EXAMPLE

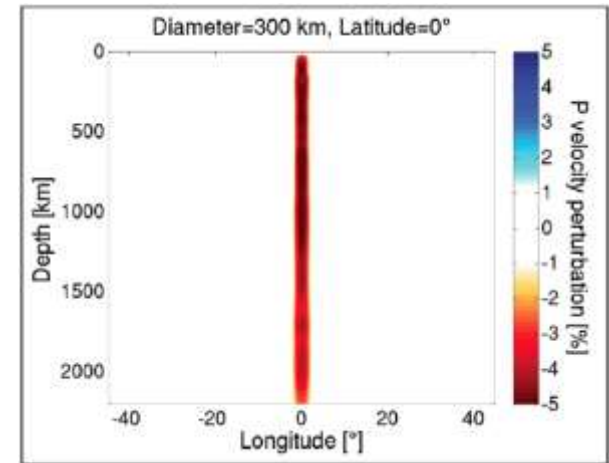
Input model – an idealised plume



reconstructed model
using only P wave traveltimes



reconstructed model
complete seismograms



Pros

- Very few mathematical approximations. [Less approximation artefacts]
- Works for realistically heterogeneous Earth models. [Sharp velocity variations of >10 %]
- Exploitation of complete seismograms. [Naturally combine body and surface wave tomography. Improve resolution, given the right data.]

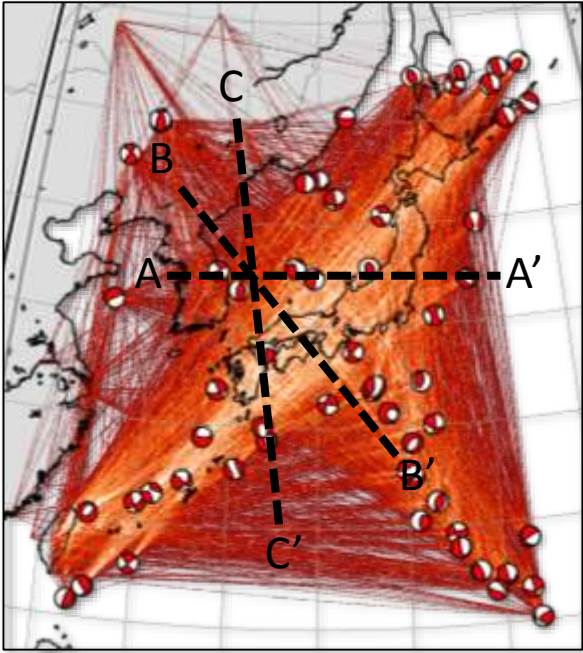
Cons

- Algorithmically complex.
- High computational requirements [due to fully numerical wave propagation].
- Use less earthquakes.

REAL-WORLD EXAMPLES

Amplitudes, crust/mantle resolution, and uncertainties

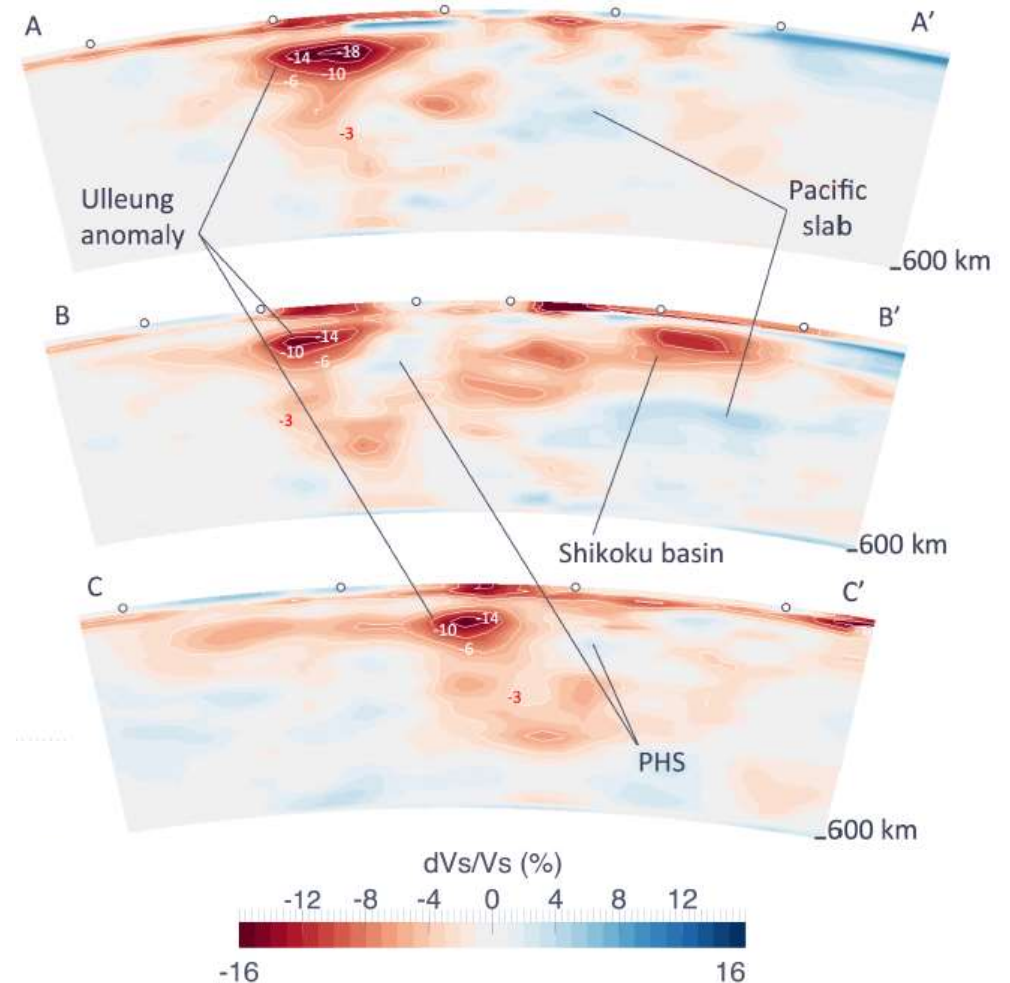
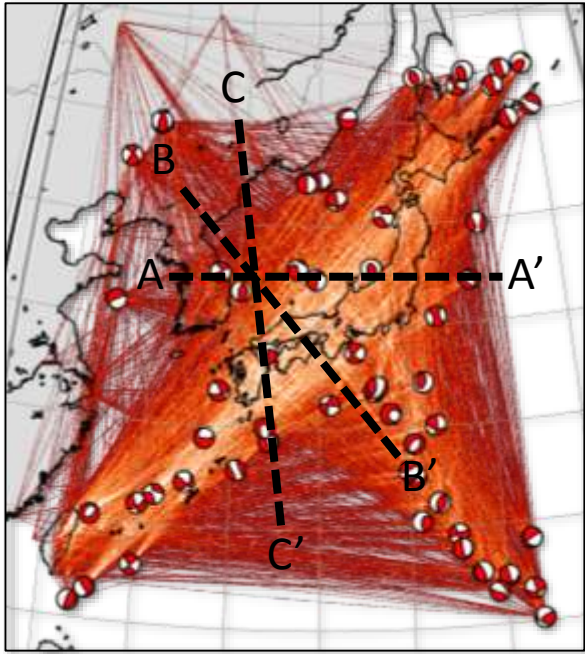
THE JAPANESE ISLANDS REGION



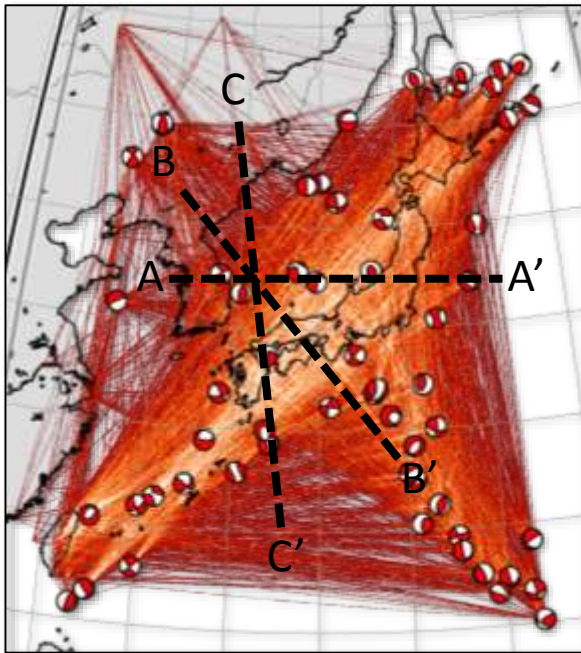
Data

- 58 earthquakes, >150 stations
- body waves, surface waves, ...
- periods: **15 – 150 s**

THE JAPANESE ISLANDS REGION



THE JAPANESE ISLANDS REGION

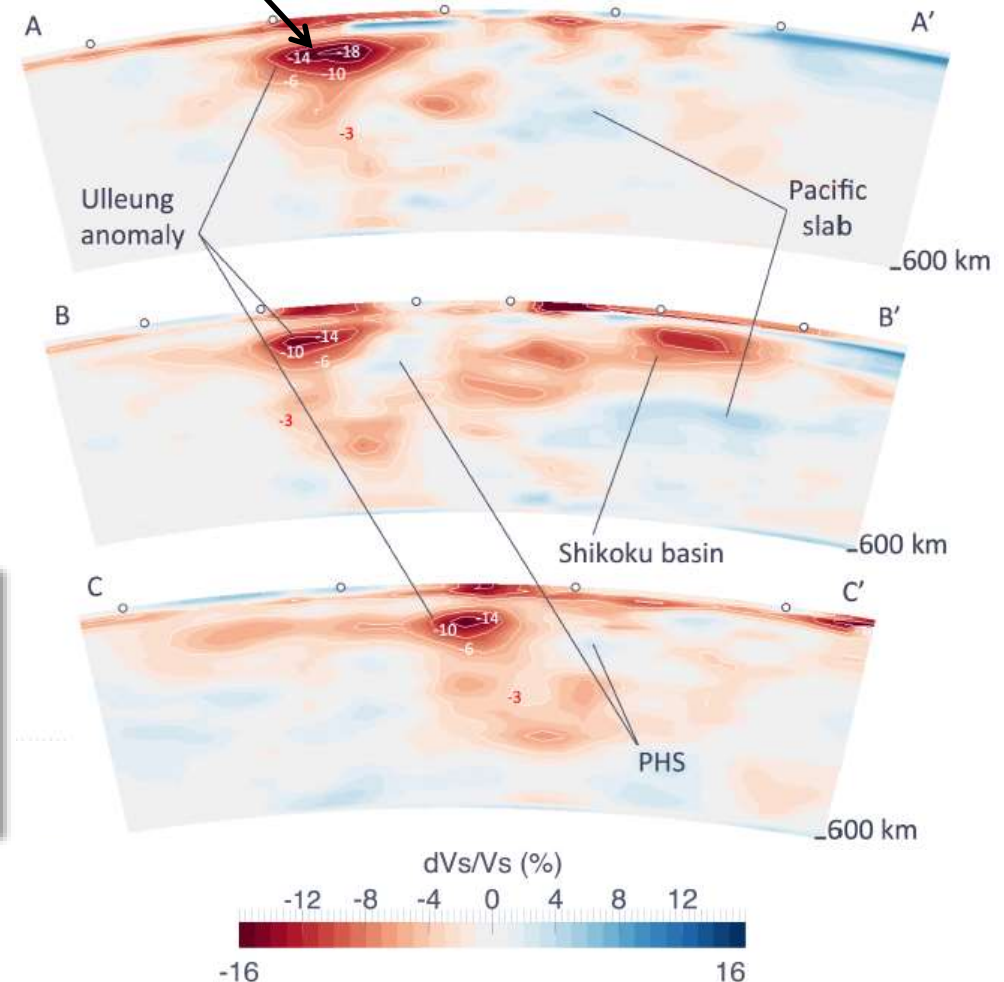


$$\delta \ln v_s = -19\%$$

[$v_s = 3.55$ km/s at 150 km depth]



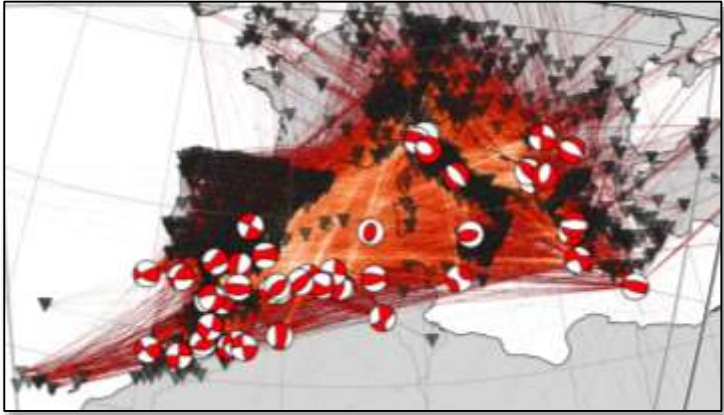
Ulleung Island



Earth is 'slower' than seen
with ray tomography

[as predicted, e.g. Wielandt 1987, Igel &
Gudmundsson 1996, Malcolm & Trampert 2011]

THE WESTERN MEDITERRANEAN

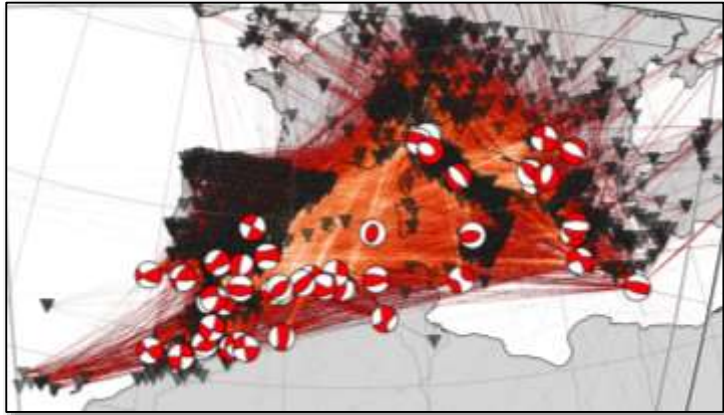


surface wave ray coverage

Data

- 52 earthquakes, >1000 stations
- body waves, surface waves, ...
- periods: **10 – 150 s**
- **6 – 90** propagation wavelengths

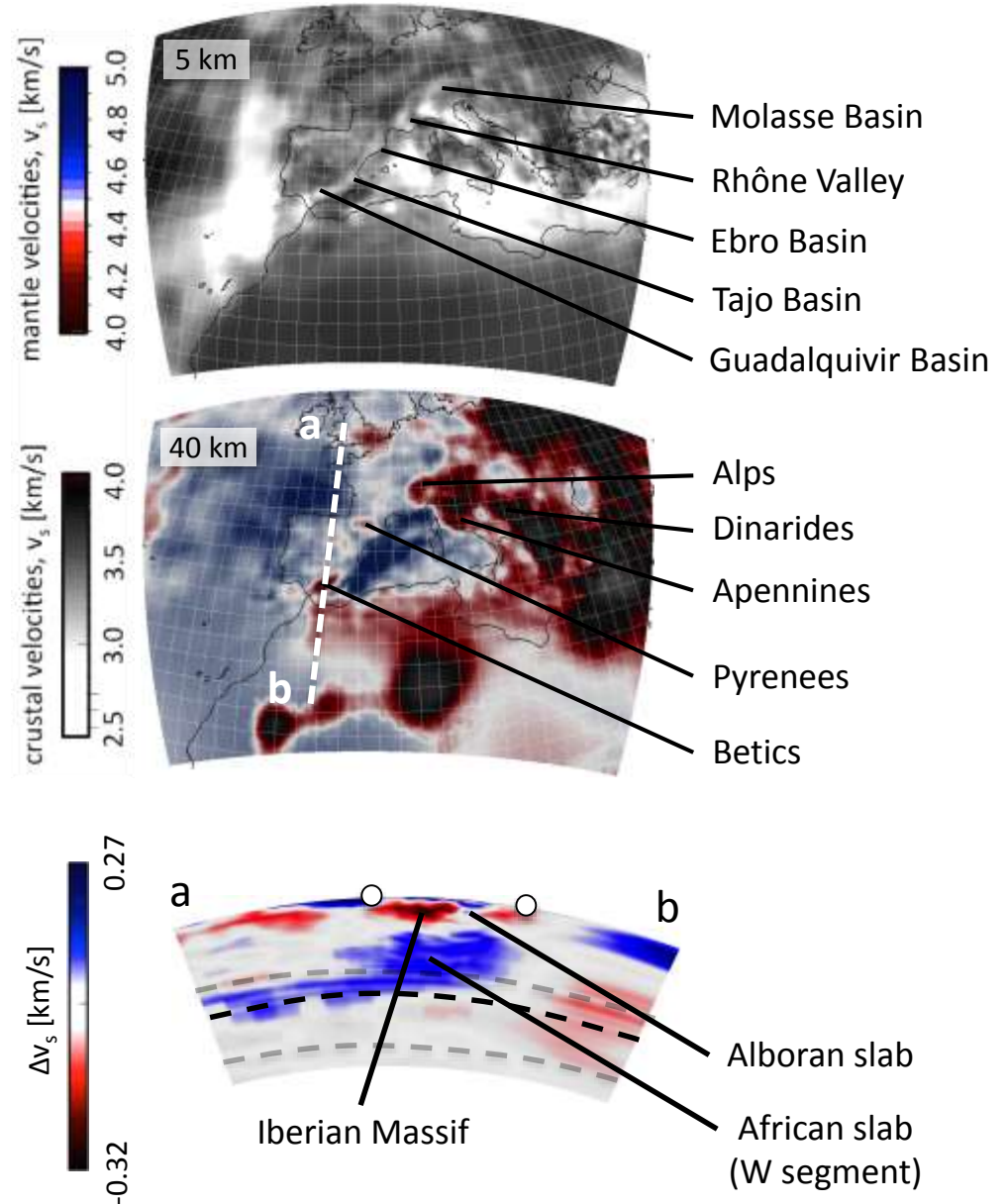
THE WESTERN MEDITERRANEAN



surface wave ray coverage

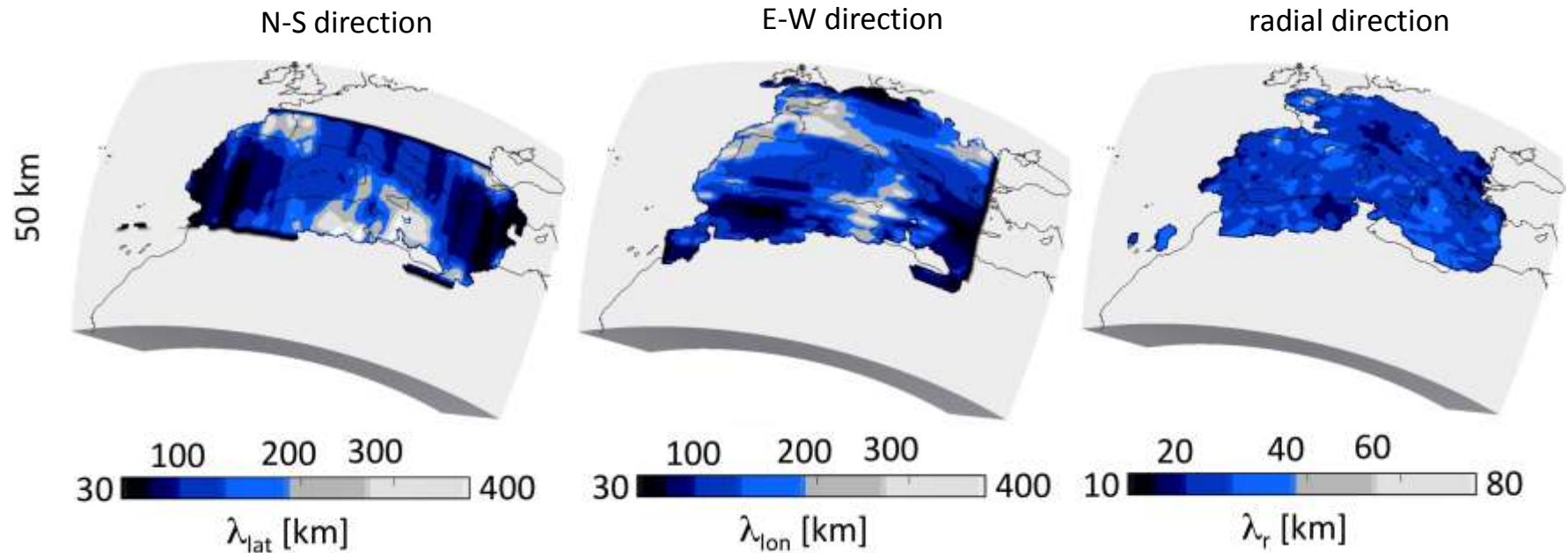
joint resolution of crust and mantle

[‘see’ relation between deep-Earth
processes and surface expressions]



Resolution analysis by random probing:

- Probing the resolution matrix with random test models [second-order adjoints]
- Direction- and position-dependent resolution lengths [point-spread function width]



efficient resolution analysis tools

[quantitative analysis instead of synthetic inverse crimes]

Intermediate take-home messages

- FWI on regional scales: It essentially works.
- Discovery of very low velocity regions. [Earth is more heterogeneous than we thought.]
 - Need to go beyond purely thermal interpretation of the model.
- Joint resolution of crustal and mantle structure
 - Direct view of relation between mantle structure and its surface imprint.
- Efficient resolution analysis tools are available.
 - Resolution is more heterogeneous than the Earth itself.

CHALLENGES

ARE OUR COMPUTERS BIG ENOUGH, OR WILL THEY SOON BE?

Compressional waves
propagate through the
whole Earth at

min. period: ≈ 1 s

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Compressional waves
propagate through the
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min. period: ≈ 1 s

Today

Global full-waveform
inversion

min. period: ≈ 20 s

Today's computing power is at least $20^5 = \mathbf{3.2 \text{ million}}$ times too small!

ARE OUR COMPUTERS BIG ENOUGH, OR WILL THEY SOON BE?

Compressional waves
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Global full-waveform
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min. period: ≈ 20 s

in ≈ 50 years

Tomography based on fully
numerical wave propagation

min. period: ≈ 1 s

Provided that:

Moore's law continues to hold.

We can handle computers that are 3.2 million times bigger.

We can write code to harness such resources.

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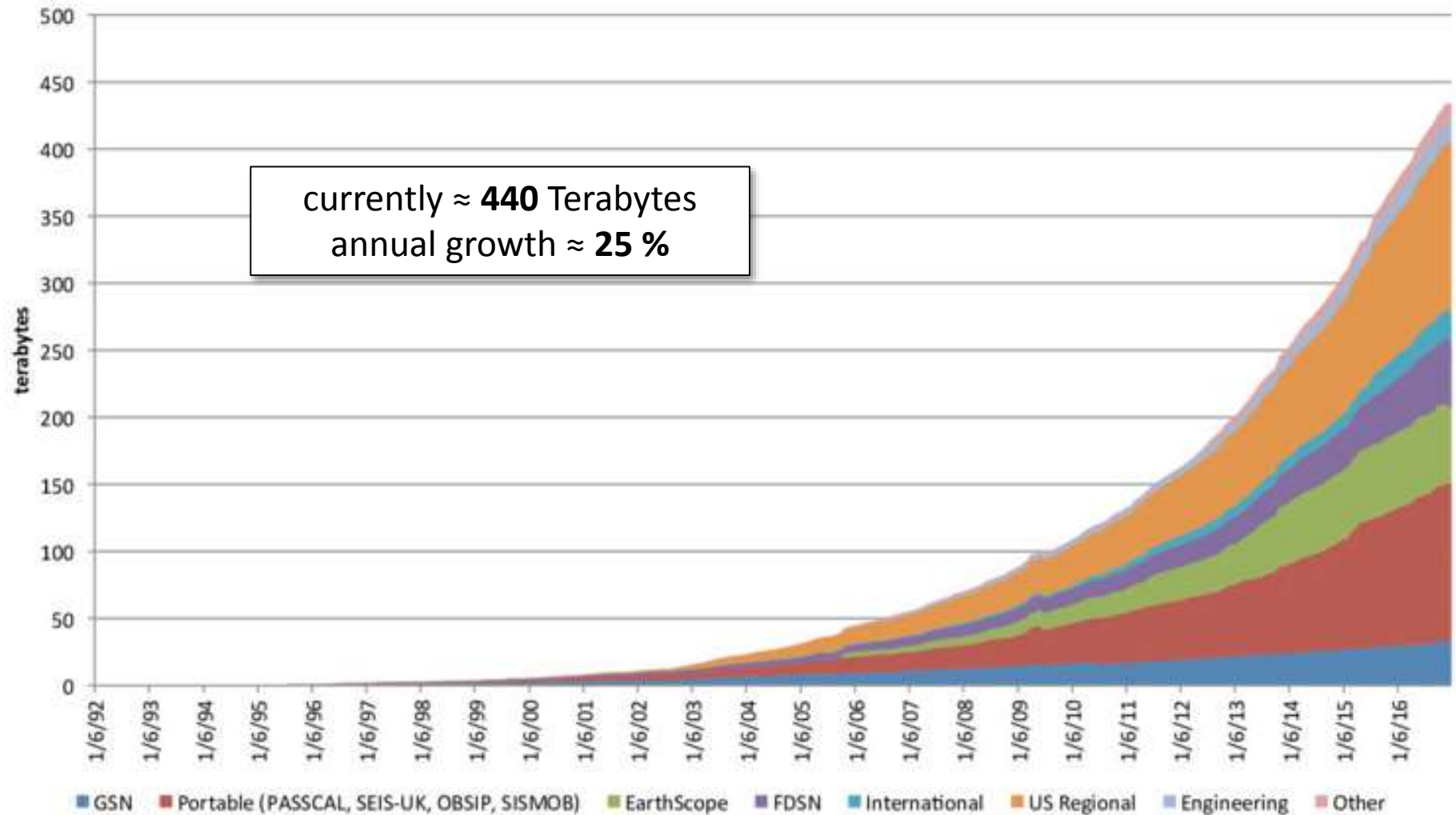
We can handle computers that are 3.2 million times bigger.

We can write code to harness such resources.

Computer power alone will not solve the problem.

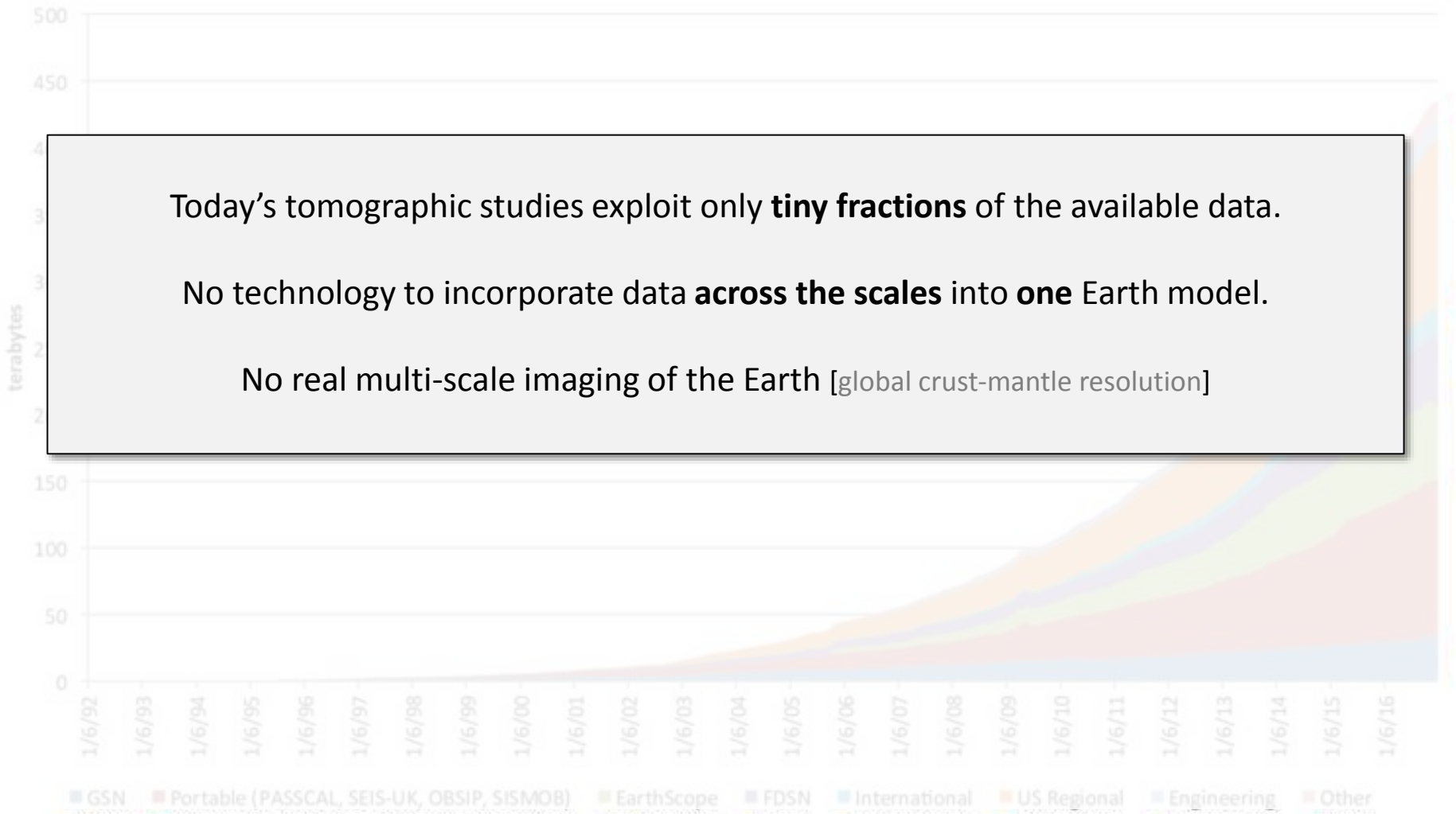
THE DATA FLOOD

IRIS Data Archive as of 1 Feb 2017



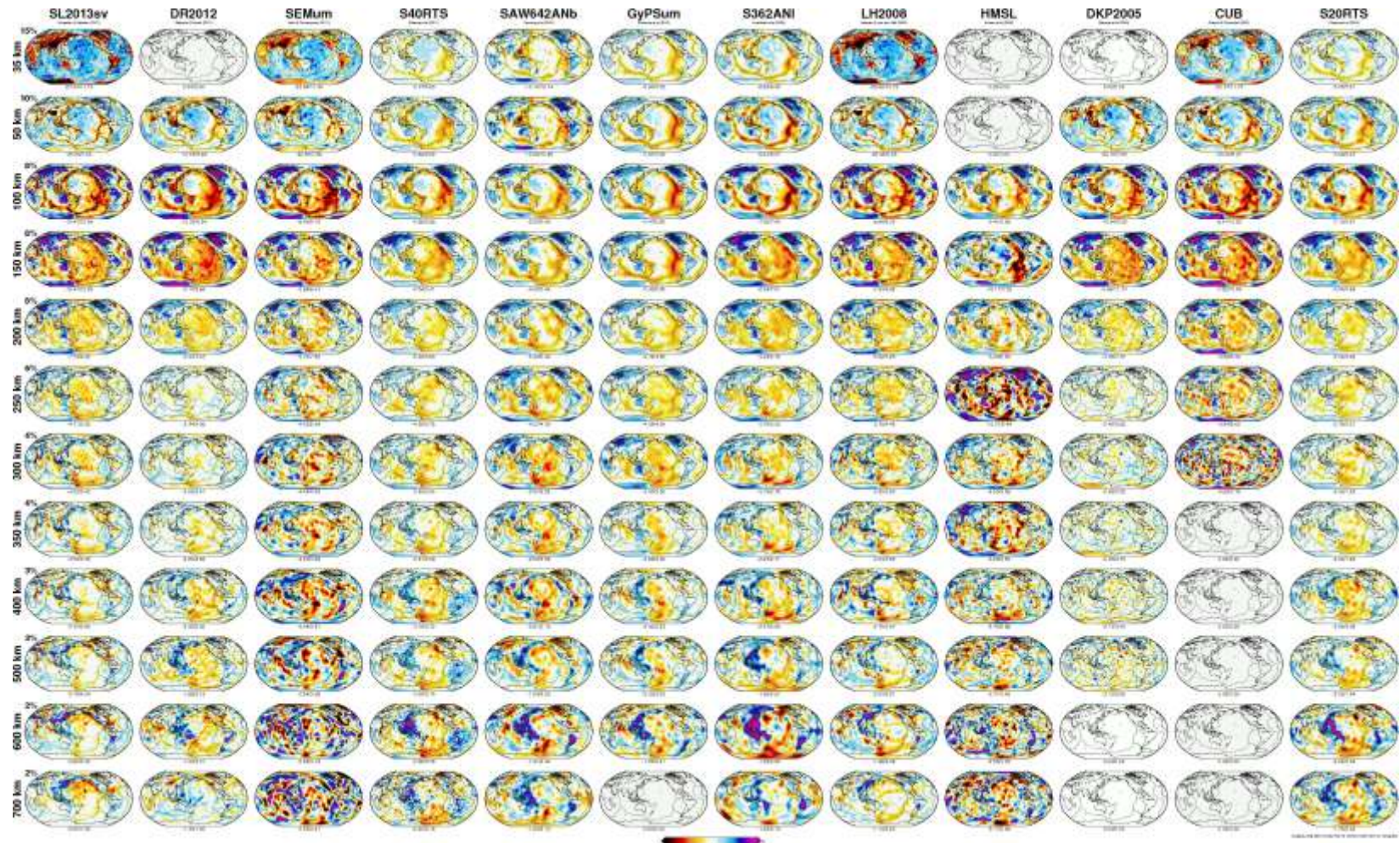
THE DATA FLOOD

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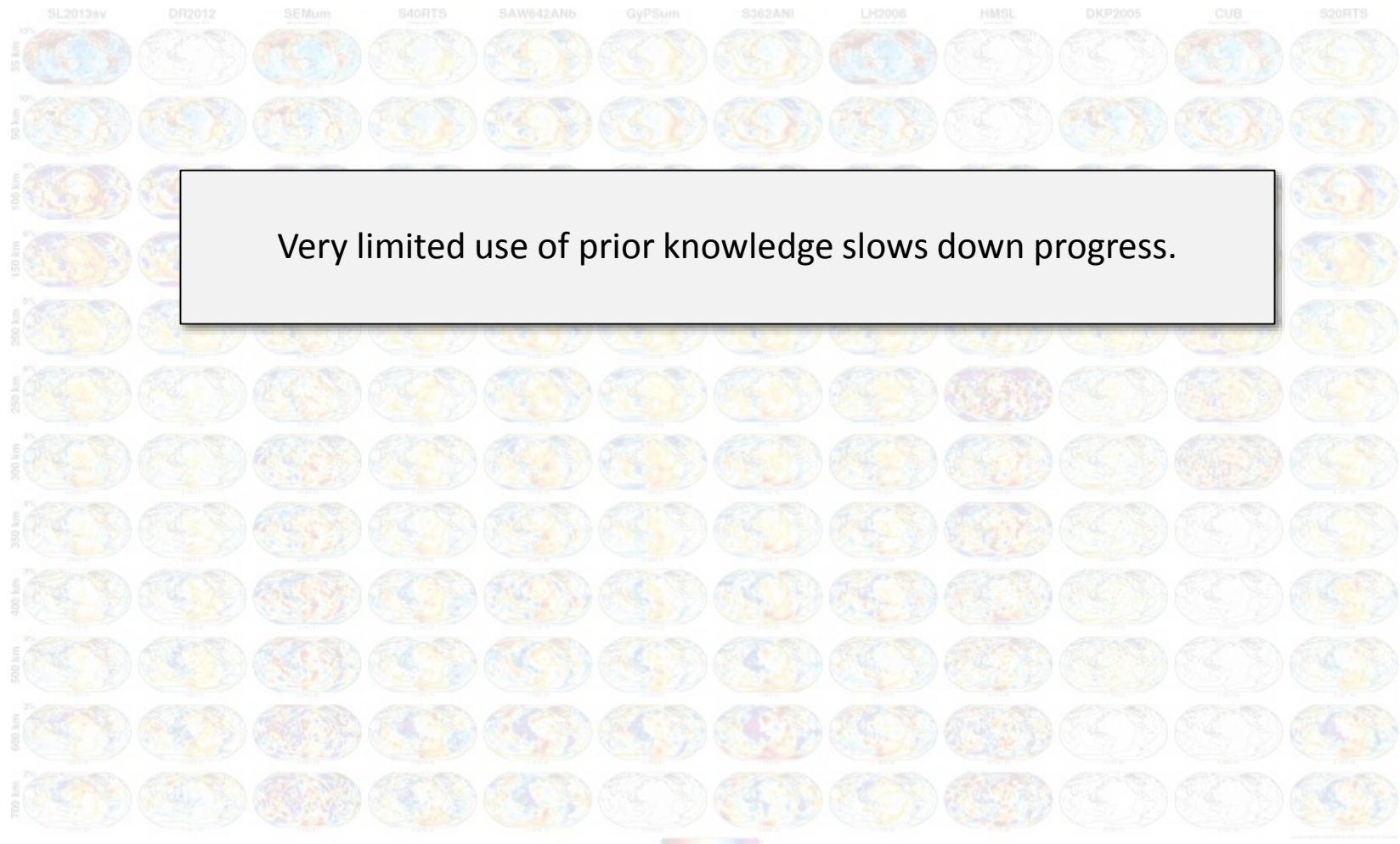
OUR MODE OF OPERATION

Comparison of 12 recent tomographic images of the same object [compiled by Andrew Schaeffer]



OUR MODE OF OPERATION

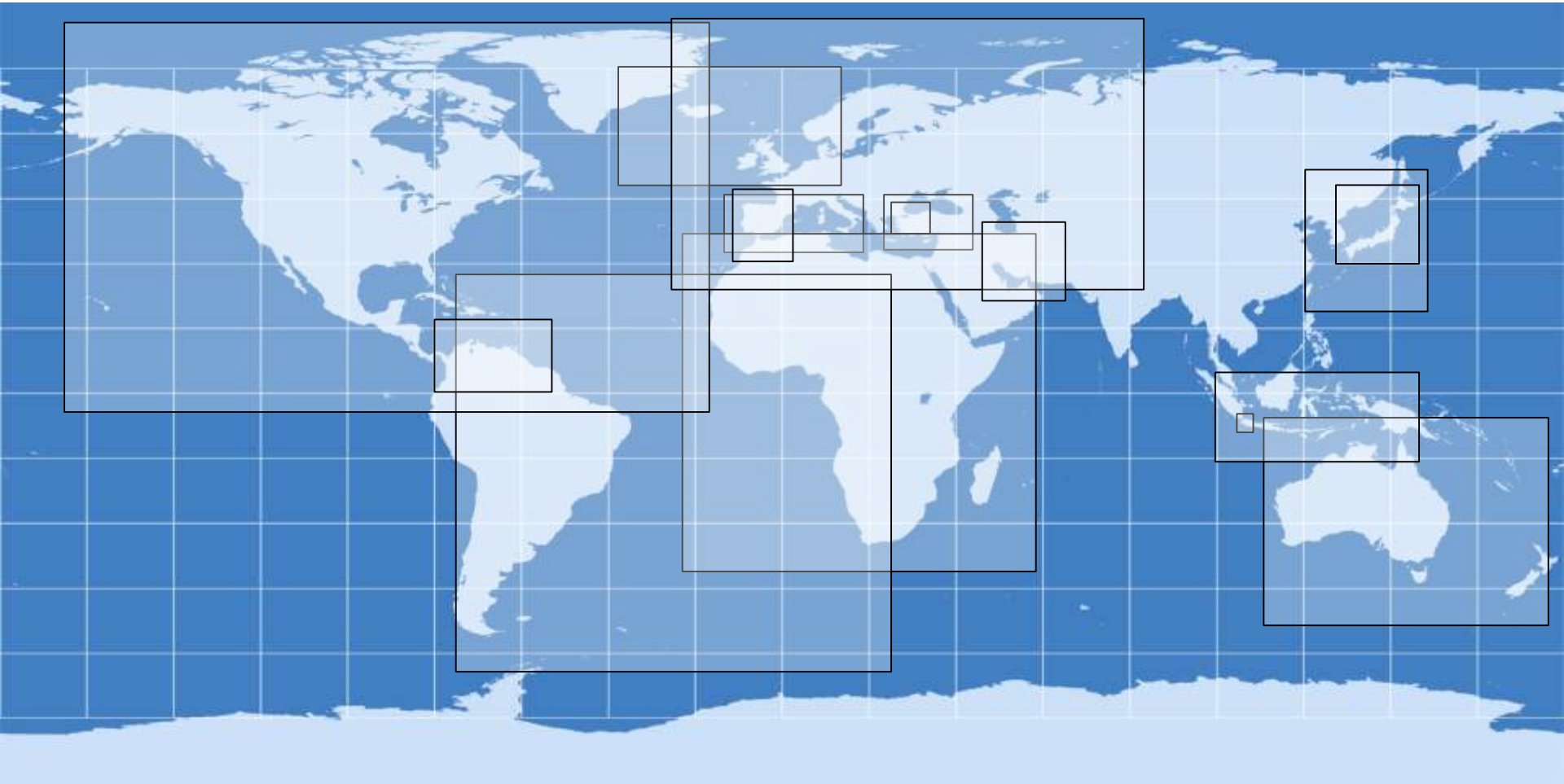
Comparison of 12 recent tomographic images of the same object [compiled by Andrew Schaeffer]



THE COLLABORATIVE SEISMIC EARTH MODEL

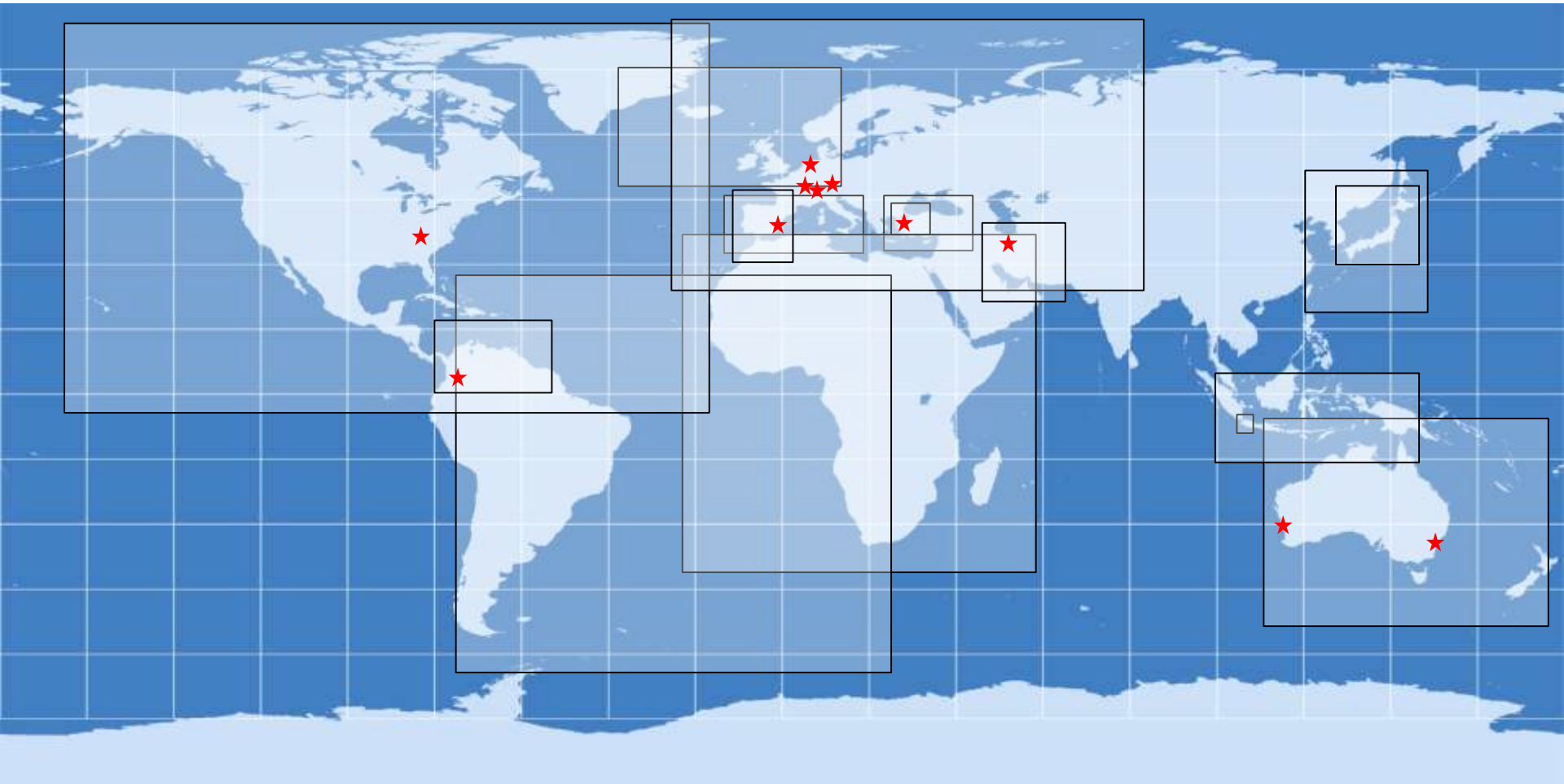
- Evolutionary multi-scale model.
- Successive regional refinements [e.g. when new data become available].
- Contributed by different researchers.
- Consistent with each other and with global Earth structure.
- Community-driven “divide and conquer”.

Overview of current subregions



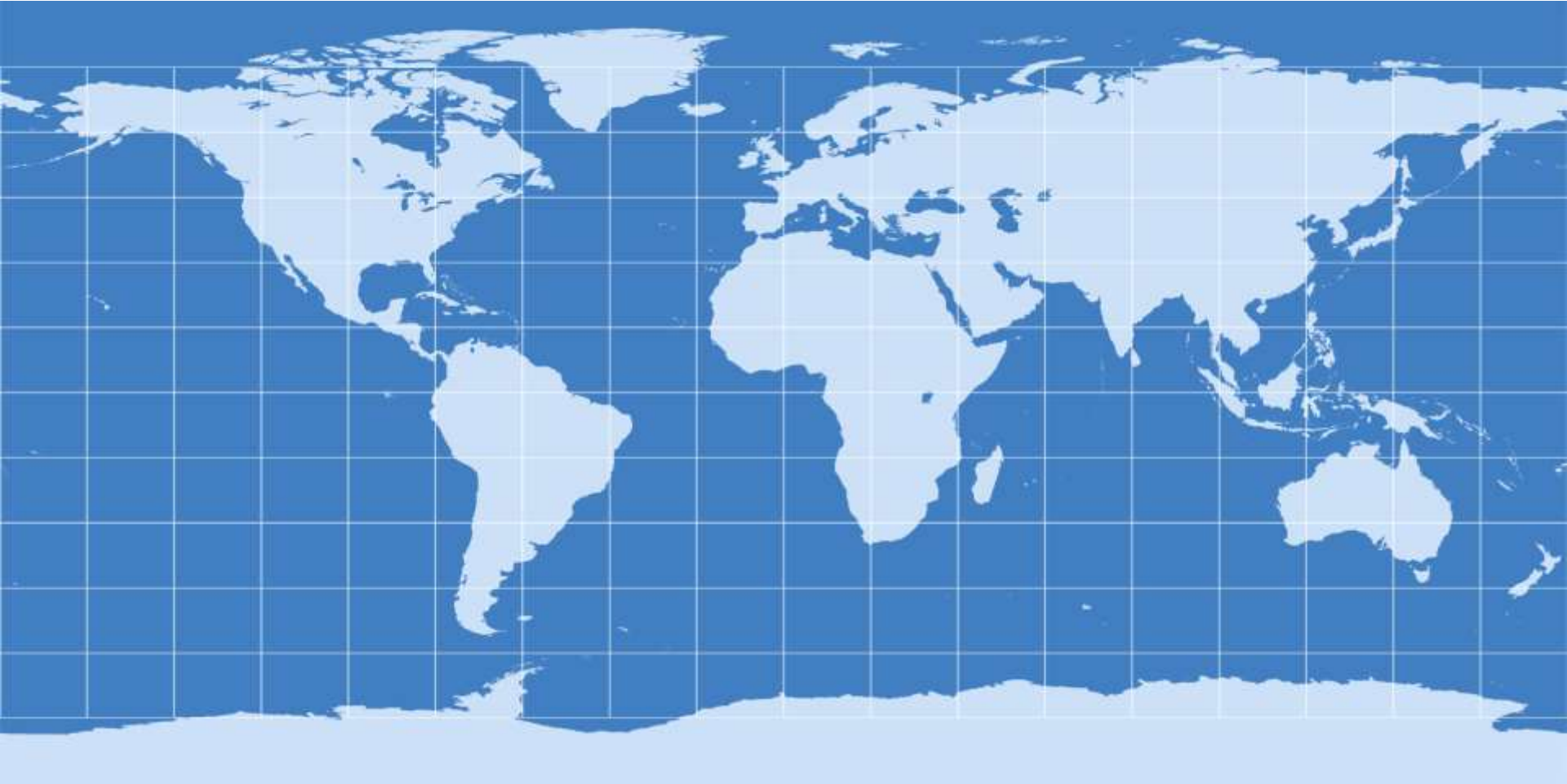
Afanasiev et al., *Geophys. J. Int.*, 2016.

Overview of current subregions



★ collaborators

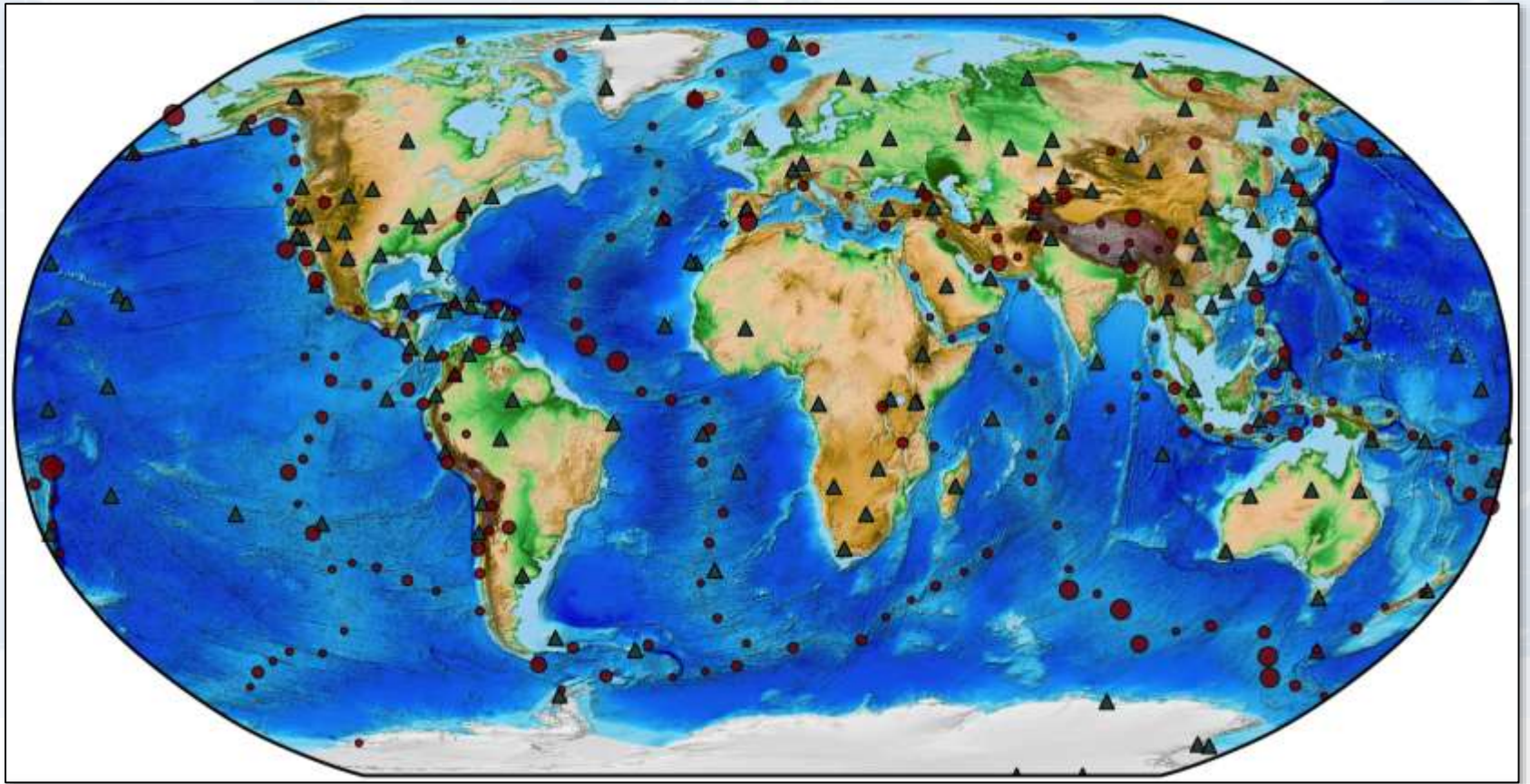
The global subregion



Afanasiev et al., *Geophys. J. Int.*, 2016.

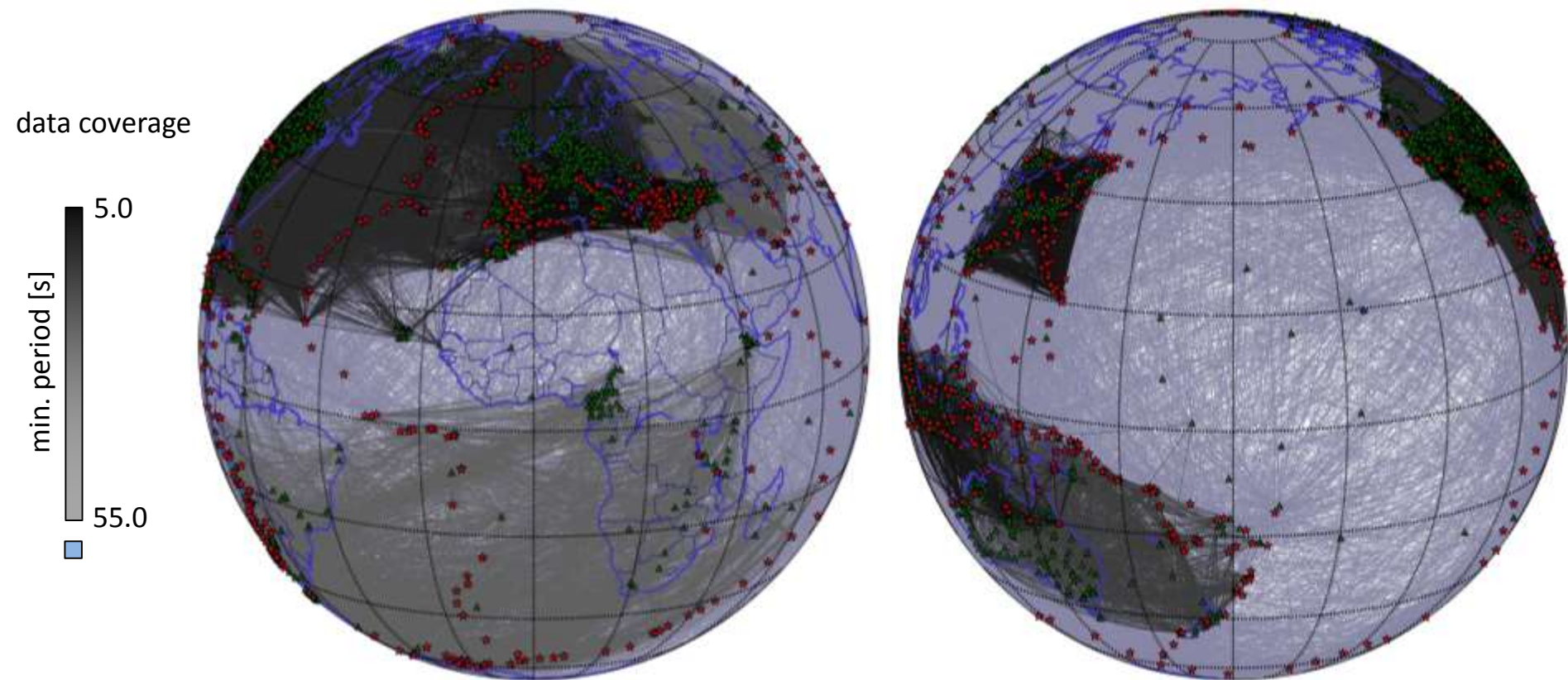
The global subregion

- Ensure consistency of regional updates with global dataset.



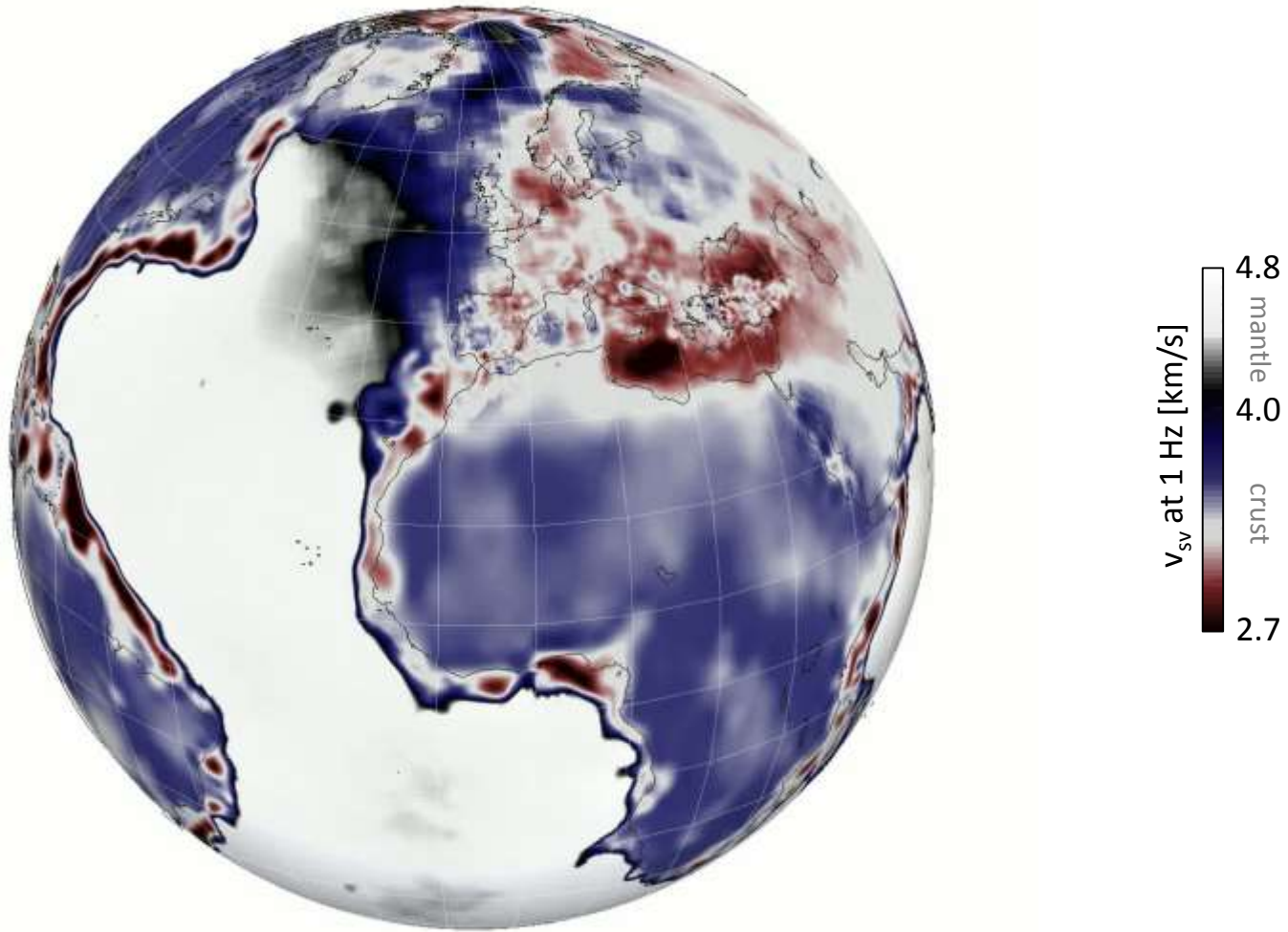
Afanasiev et al., *Geophys. J. Int.*, 2016.

CURRENT DATA COVERAGE



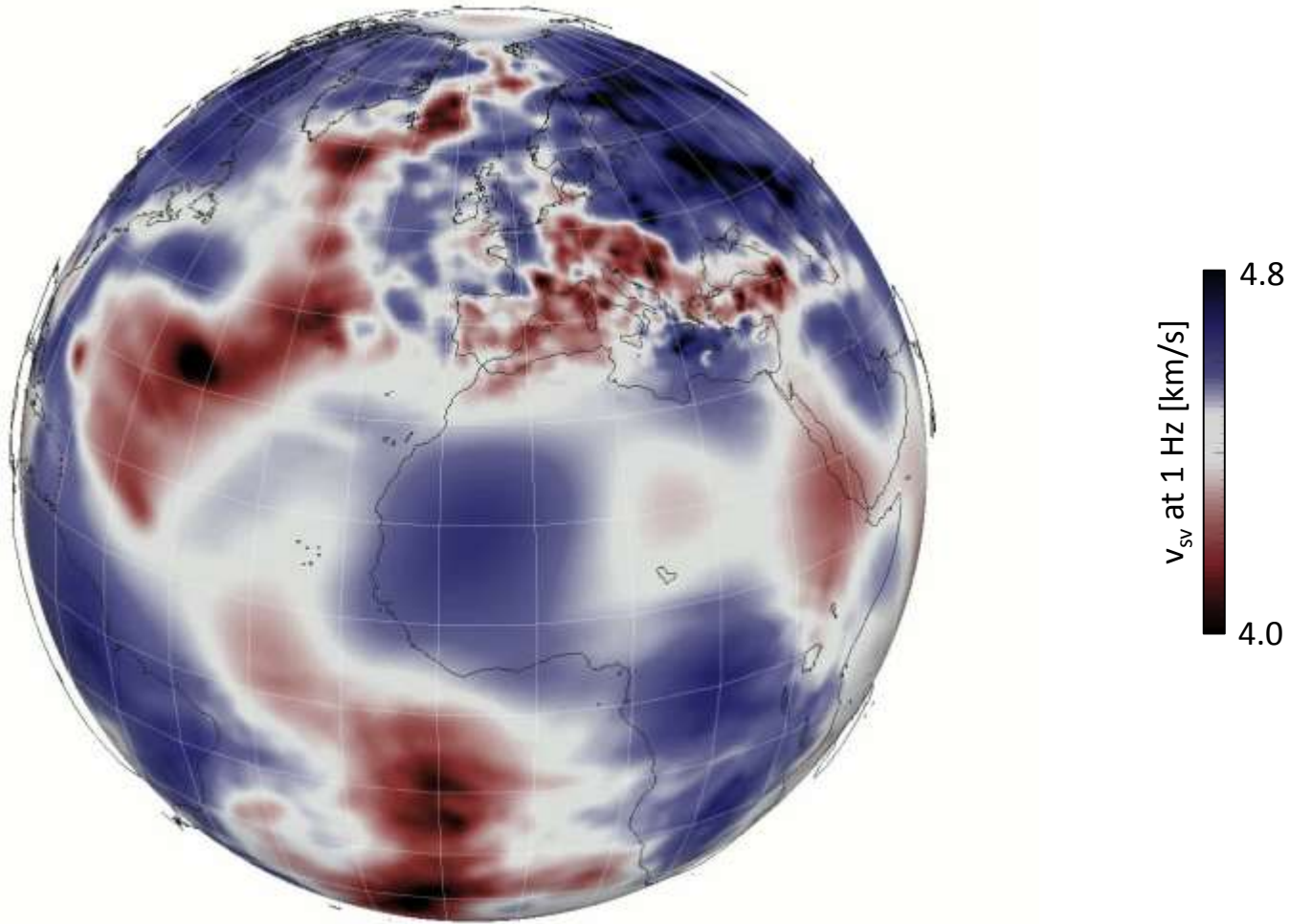
GENERATION 1

shear velocity at 15 km depth

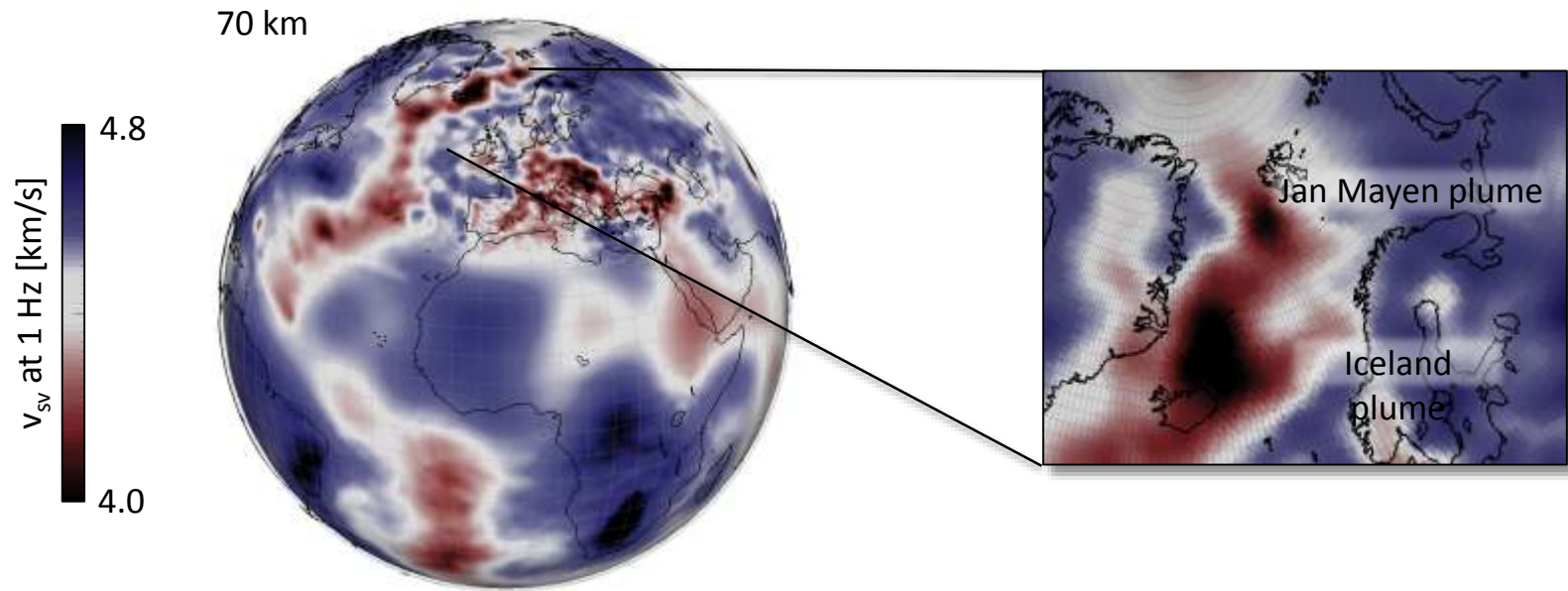


GENERATION 1

shear velocity at 100 km depth



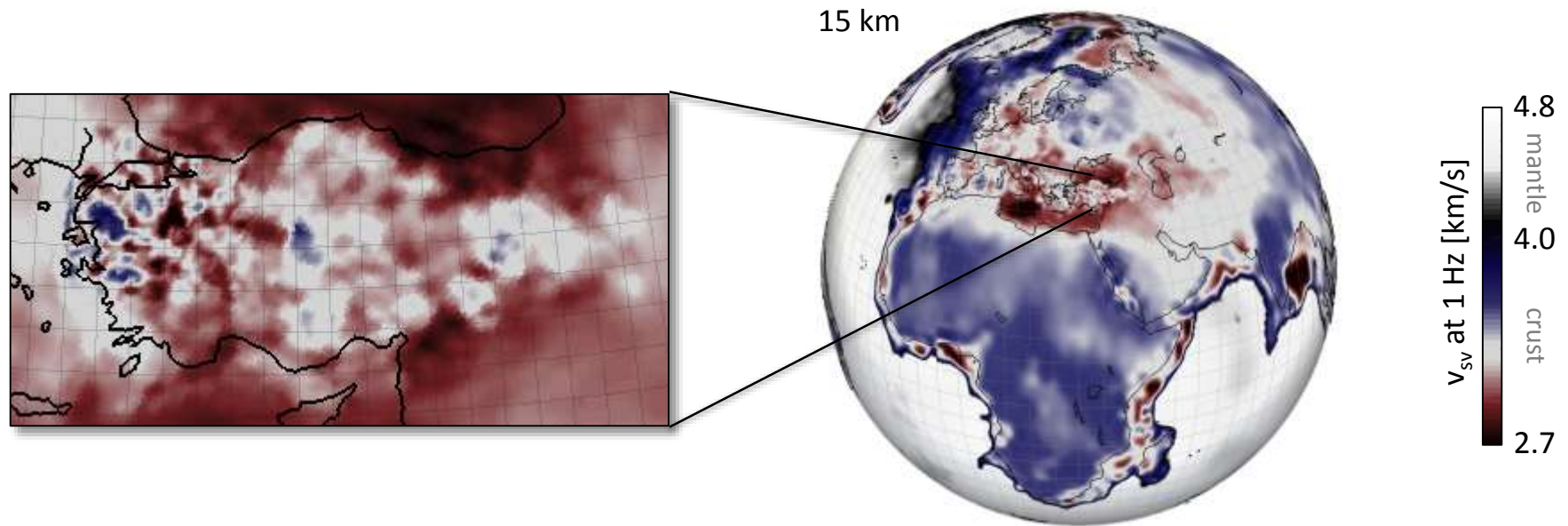
GENERATION 1



Understanding the Earth's dynamics and evolution

- Hot convective upwellings from 100's – 1000's km depth.
- Key to understand: volcanism, heat budget and evolution of the Earth.

GENERATION 1

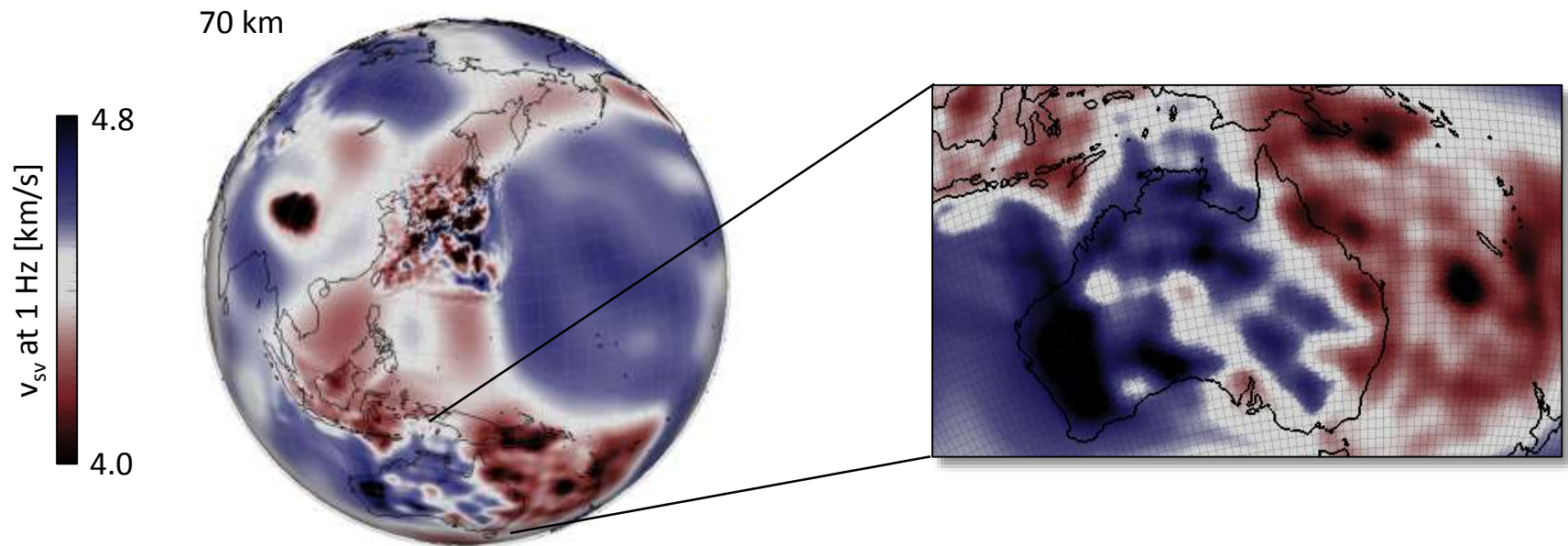


Towards earthquake ground motion prediction

- Anticipate ground motion caused by a given earthquake.
- Inform engineers, building codes,

Near-real-time earthquake characterisation

- Better Earth model → better information on earthquake properties.
- Key to improve tsunami early warning systems.
- Joint project with *Australian National University* and *Geoscience Australia*.



FWI: Range of optimal applicability

- High-quality data. [You trust all the wiggles that you exploit.]
- Strongly heterogeneous medium. [Δv around 10 % or more, ocean-continent boundary]

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FWI: Range of methodological overkill

- Only traveltimes are trustworthy. [insufficient data quality]
- Harmless medium. [e.g., most of the Earth below 300 km depth]

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FWI: Status quo

- Discovery of smaller and stronger heterogeneities.
- Interpretation requires incorporation of compositional effects [anomalies too big for being purely thermal].
- **Earth may be more heterogeneous than we thought.**

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- Link mantle structure and surface observables.

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- Powerful random probing techniques make uncertainties accessible.
- Quantitative resolution analysis. **Go beyond the chequerboard!**

Big challenges [not necessarily limited to FWI]:

- Computer power will remain insufficient for a long time [good, if you like using your brain].
- Data flood requires man power that an individual group does not have.
- Individualism and insufficient use of prior knowledge.
- Remaining problem: Change the community's mode of operation.

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- Remaining problem: Change the community's mode of operation.

The Collaborative Seismic Earth Model

- Vision of a community-driven, evolutionary, global multi-scale model.
- Continued and consistent refinements on all scales.
- Generation 1 exists, with currently around 10 collaborators worldwide.

Big challenges [not necessarily limited to FWI]:

- Computer power will remain insufficient for a long time [good, if you like using your brain].
- Data flood requires man power that an individual group does not have.
- Individualism and insufficient use of prior knowledge.
- Remaining problem: Change the community's mode of operation.

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Thanks for your attention!